

Update of the EIGEN time variable gravity model for precise orbit determination

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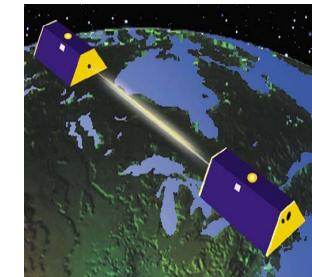
(1) CNES/GRGS, Toulouse, France

(2) Géode & Cie, Toulouse, France

(3) GET/UMR5563/OMP/GRGS, Toulouse, France

GRACE (L-1B “Version2” data)

- K-Band Range-Rate data ($\sigma_{\text{apriori}} = .1 \mu\text{m}$)
- Accelerometer / attitude
- GPS data (1-day arcs, $\sigma_{\text{code}} = .8 \text{ m}$, $\sigma_{\text{phase}} = 20 \text{ mm / 30s resolution}$)
(previously: $\sigma_{2002-2003} = 8 \text{ mm/30 s}$, $\sigma_{2003-2013} = 20 \text{ mm/300 s}$, $\sigma_{2013-2016} = 8 \text{ mm/30 s}$)



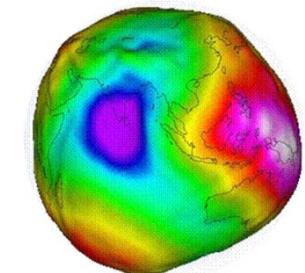
SLR

- Lageos1/2 data (10-day arcs, $\sigma_{\text{apriori}} = 6 \text{ mm}$)
- Starlette/Stella data (5-day arcs, $\sigma_{\text{apriori}} = 10 \text{ mm}$)



Physical parameters present in the normal equations

- Gravity spherical harmonic coefficients complete to **degree and order 80** (truncated to **30** for LAGEOS and **40** for GPS data)
- Ocean tides s. h. coefficients for 14 tidal waves with maximum degree/order ≤ 30 (not used yet)



Models used : v2 → v3

Dynamical models

Gravity	<i>EIGEN-6S2</i> → <i>EIGEN-GRGS.RL03-v2.MEAN-FIELD</i>
Ocean tide	<i>FES2012 (Legos)</i> → <i>FES2014 (Legos)</i>
Atmosphere	<i>3-D ECMWF ERA-interim pressure grids / 3hrs</i>
Ocean mass model	<i>TUGO (Legos) / 3hrs</i>
Atmospheric tides	→ <i>Not necessary because of the 3hrs dealiasing time sampling</i>
3 rd body	<i>Sun, Moon, 6 planets (DE405)</i>
Solid Earth tides	<i>IERS Conventions 2010</i>
Pole tides	<i>IERS Conventions 2010</i>
Non gravitational	<i>Accelerometer data (+biases and scale factors)</i>

Geometrical models

SLR stations	<i>ITRF2008 coordinates</i> → <i>ITRF2014 coordinates</i>
GPS	<i>IGS Repro-1 orbits & clocks</i> → <i>IGS Repro-2 orbits & clocks</i>

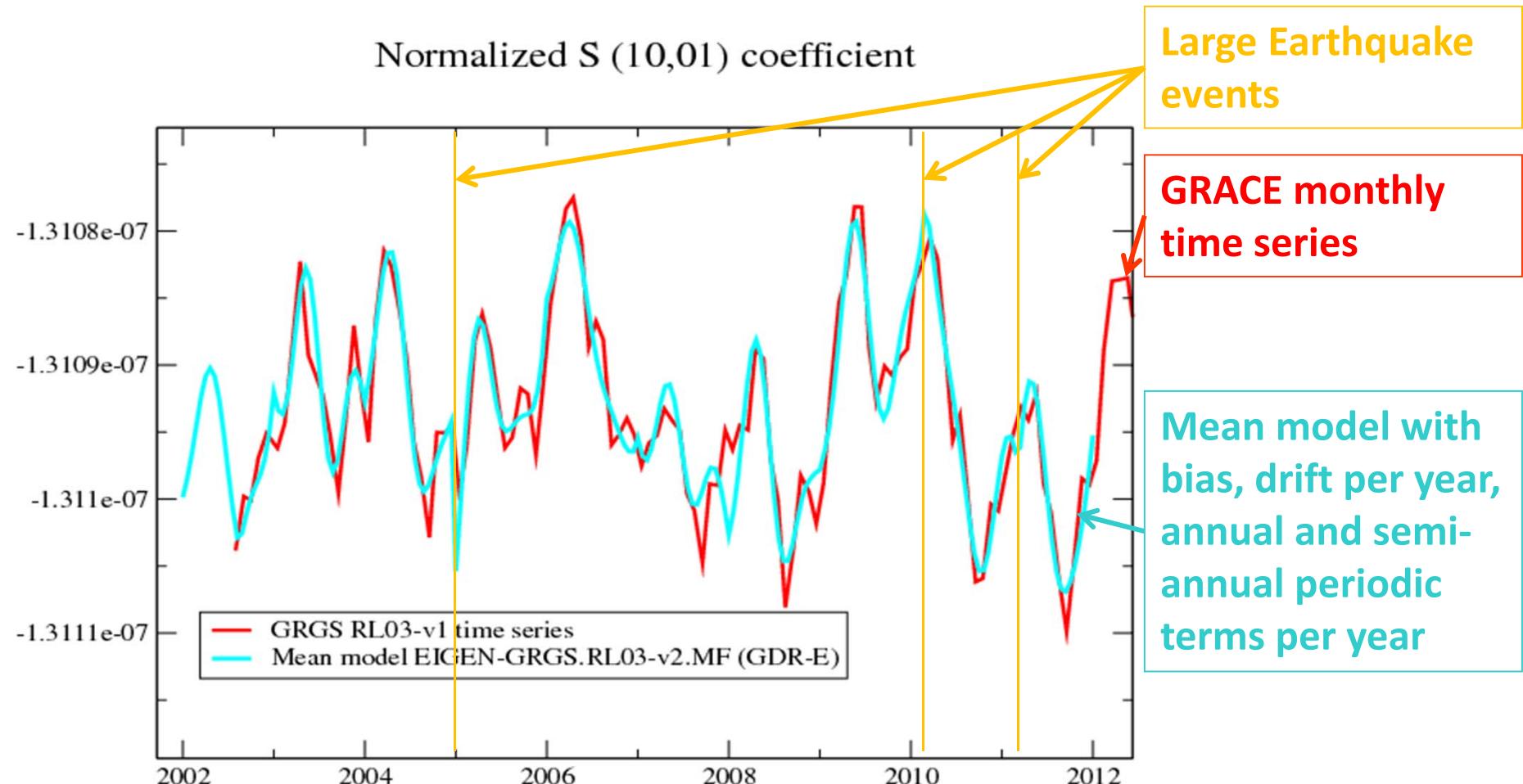
Other models

Hydrology	Taken into account by the a priori gravity field
Glacial Isostatic Adjustment	

❖ Mean models are now generated from time series

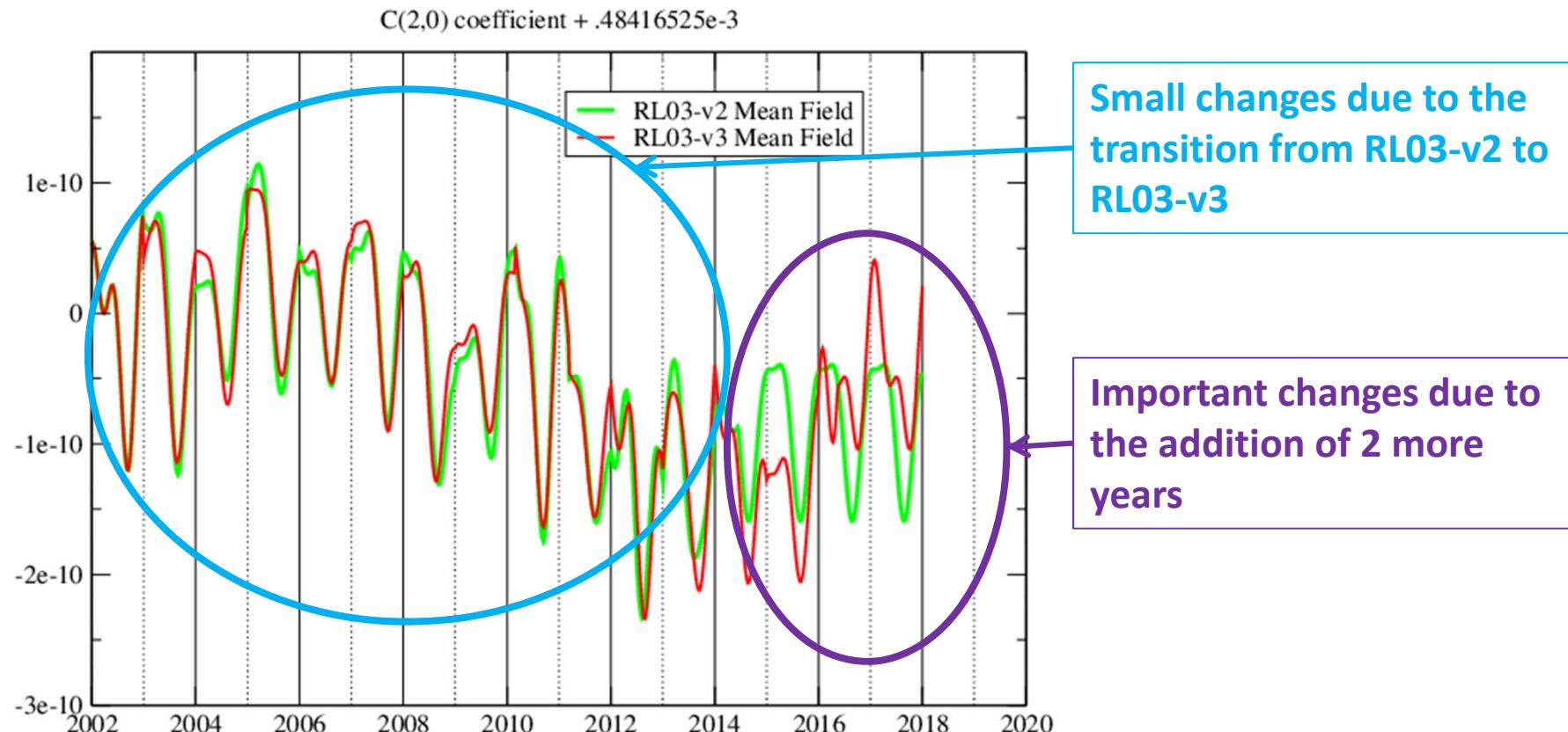
- Fitting each series of monthly spherical harmonic coefficients by a set of 6 parameters :
 - Yearly bias and slope : piecewise linear function except in case of ...
 - Jumps caused by big earthquakes (so far : Sumatra/2005.0, Concepcion/2010.2 and Tohoku/2011.2)
 - Annual and semi-annual sine/cosine functions (with continuity constraints at hinge epochs)
- It means 600 000 coefficients for a 80x80 spherical harmonic model which better match with GRACE monthly models
- Used for operational computation (i.e. altimetric orbit processing) or TRF processing (i.e. ITRF2014)

➤ Example for one spherical harmonic coefficient:



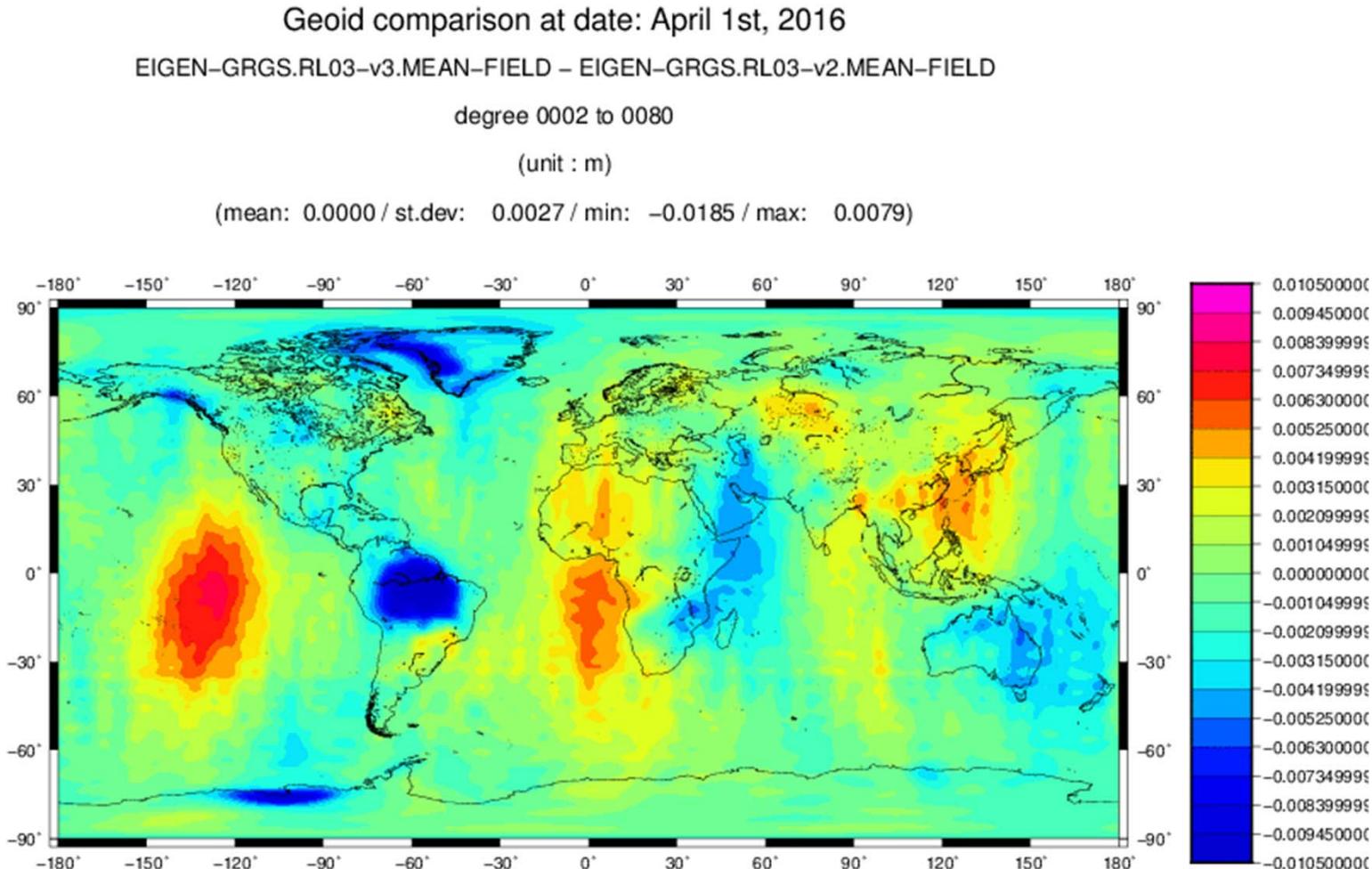
Update of the mean model from -v2 to -v3

- The new mean field updates the previous one over 2 years: mid-2014 to mid-2016.
- Example for the C(2,0) spherical harmonic coefficient:



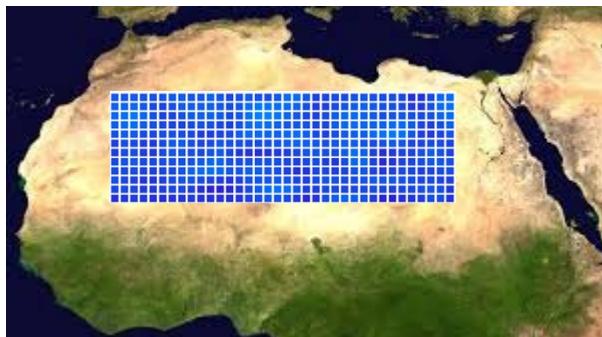
Update of the mean model from -v2 to -v3

- Extrapolation vs. real data after 2 years: difference between mean-field –v2 and mean field –v3 at mid-2016

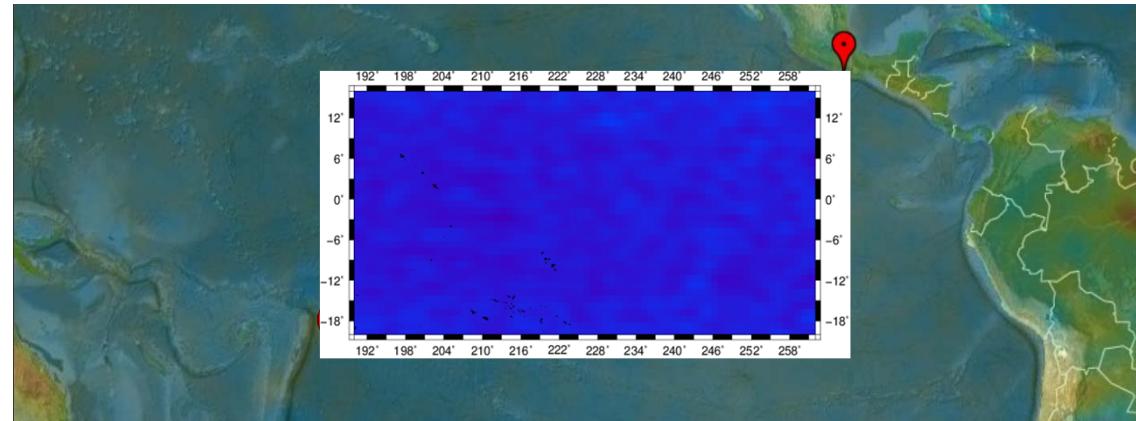


RL03/RL04 evaluation

Noise assessment can be made in areas with no or very little mass variations:
Sahara and Gobi deserts, East Antarctica, South and Equatorial Pacific



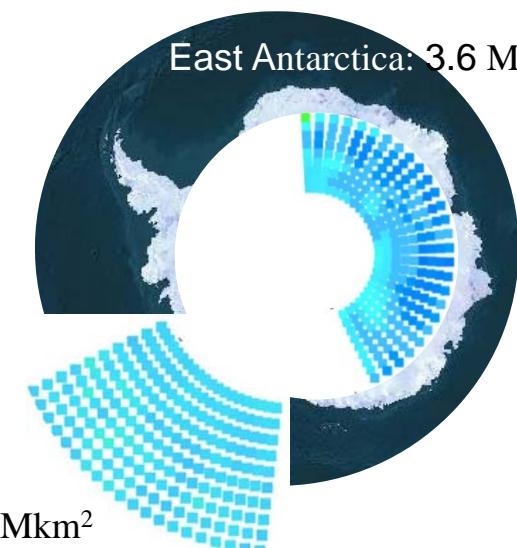
Sahara desert: 2.2 Mkm^2



Equatorial Pacific: 31.6 Mkm^2



Gobi desert: 1.6 Mkm^2



East Antarctica: 3.6 Mkm^2

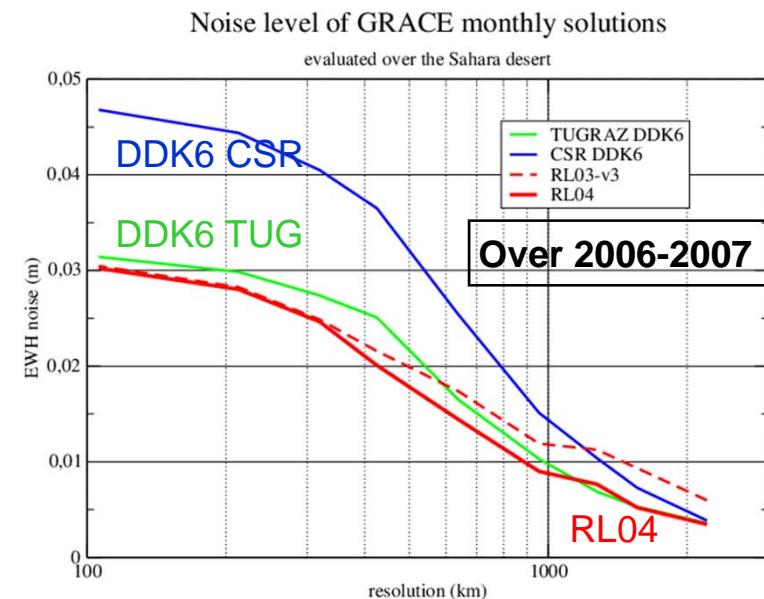
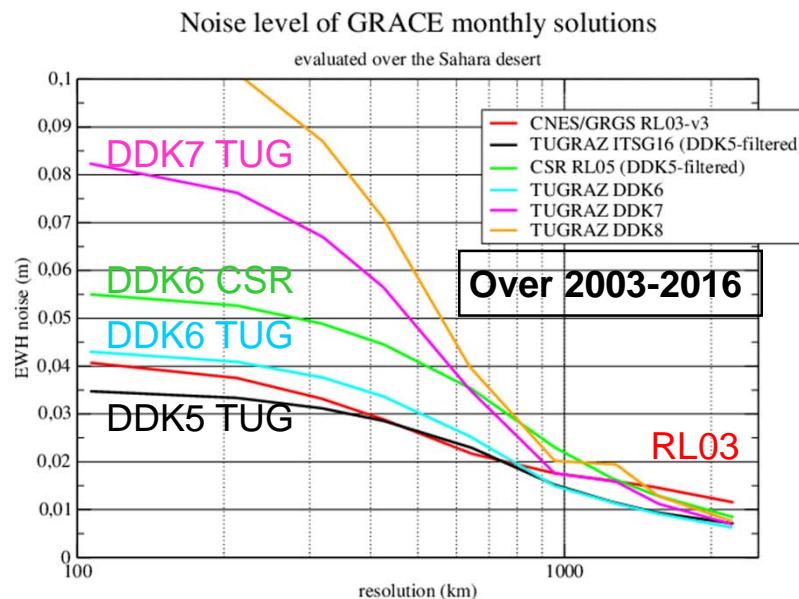
South Pacific 6.7 Mkm^2

Noise assessment over the Sahara

The Sahara desert shows very little hydrological variations. We have delimited a rectangular zone of 2.2 Mkm² where almost no gravity variation is suspected. (except a small depletion of 1.3 mm/yr in South Libya).

It is hence well dedicated to control the quality of gravity field variation models. The surface is first divided in 2 deg.*2 deg. blocs (\Leftrightarrow degree/order 90), then averaged in blocs of larger size up to 20 deg.*20 deg. Drift and annual/semi-annual variations are fitted a priori.

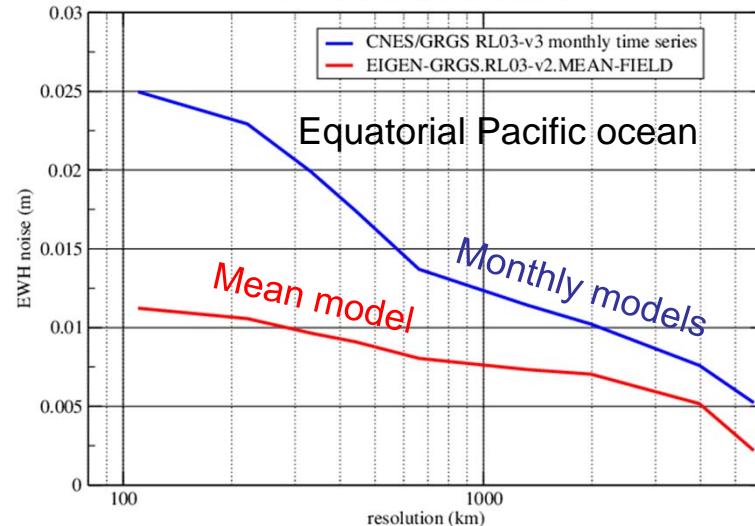
Different time-varying gravity models with various DDK filters (Kusche et al.) are compared spectrally in this way from 100 km to 2200 km.



Noise assessment over “deserts”

Noise level of GRACE monthly solutions

evaluated over the Equatorial Pacific ocean



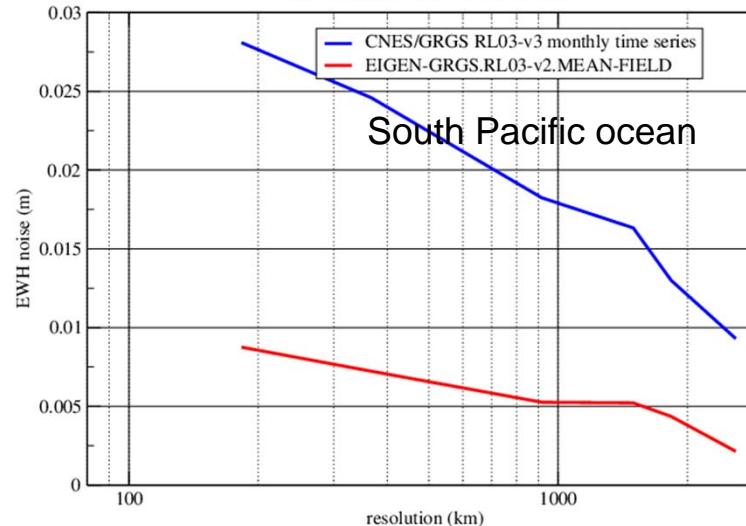
Equatorial Pacific ocean

Mean model

Monthly models

Noise level of GRACE monthly solutions

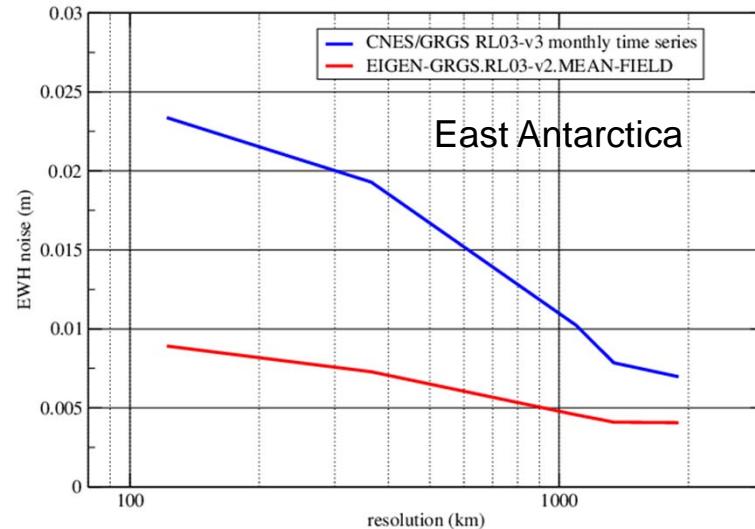
evaluated over the South Pacific ocean



South Pacific ocean

Noise level of GRACE monthly solutions

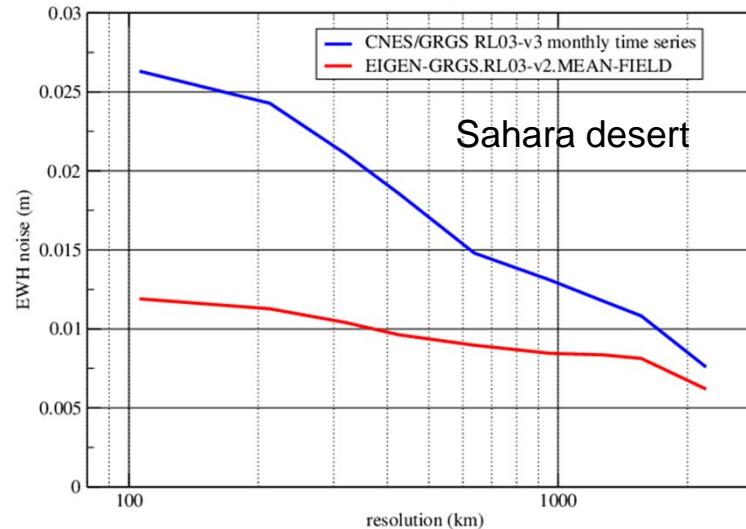
evaluated over East Antarctica



East Antarctica

Noise level of GRACE monthly solutions

evaluated over the Sahara desert

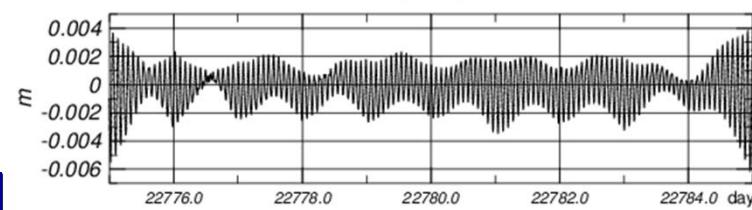
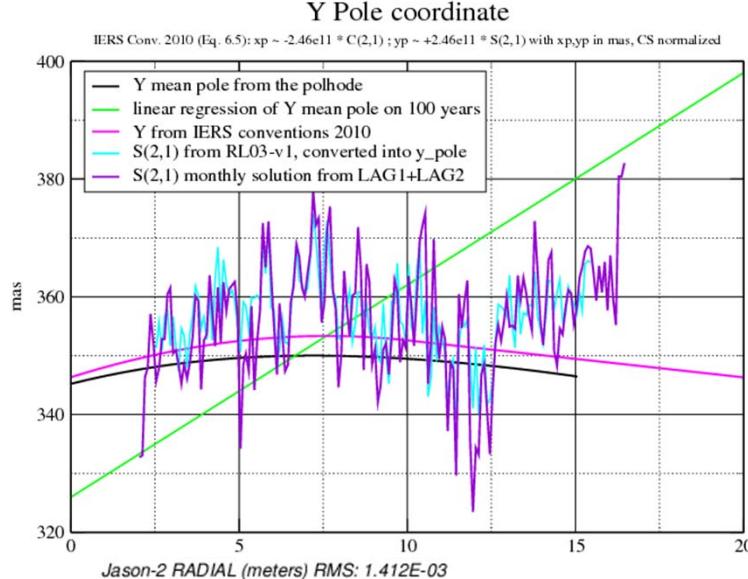
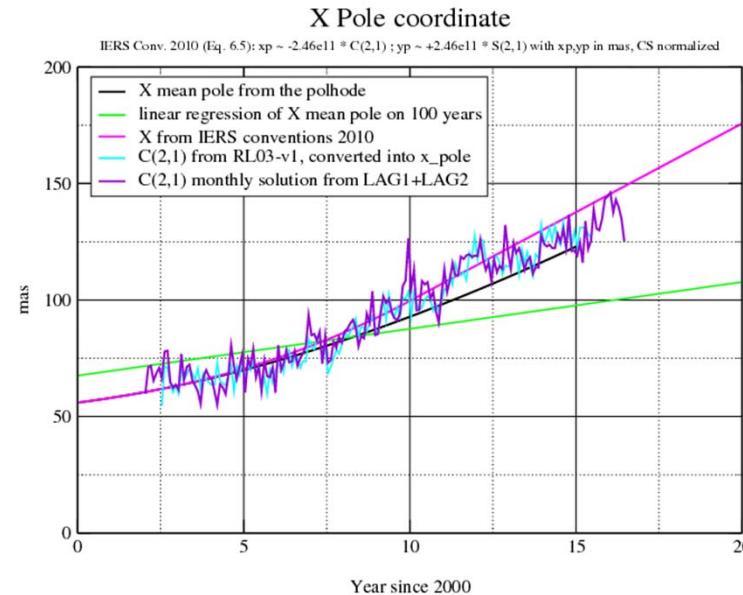


Sahara desert

Coherence gravity / mean pole model

Very important:

- When using the C(2,1)/S(2,1) values of a gravity field model, one must adopt **the same mean pole convention** as the one used for the computation of the model.
- The same goes for the **dealiasing models**.
- CNES/GRGS is using the mean pole of the **IERS2010 conventions**. If the conventions change for a **linear mean pole**, then the C(2,1)/S(2,1) coefficients of the mean gravity model will have to be adapted to this new convention.



- The new mean gravity field model based on RL03-v3 is ready to succeed RL03-v2 in the new GDR standards
- It is expanded in sets of 6 yearly coefficients (bias, slope, annual and semi-annual terms) per degree/order up to 80, and contains constant terms (from GOCE-DIR5) up to degree/order 300
- Extrapolated time-variable terms (before August 2002 and after May 2016) are based on global fits of monthly coefficients over 14 years of GRACE data
- Degree 2 time-variable terms are adjusted back to 1986 from Lageos data

- We are presently reprocessing all the GRACE data with new standards in order to produce the Release 04 of our time-variable gravity solutions
- We expect to have completed the reprocessing by the end of December 2017
- From this RL04, a new mean model up to degree 90 could be produced as early as January 2018