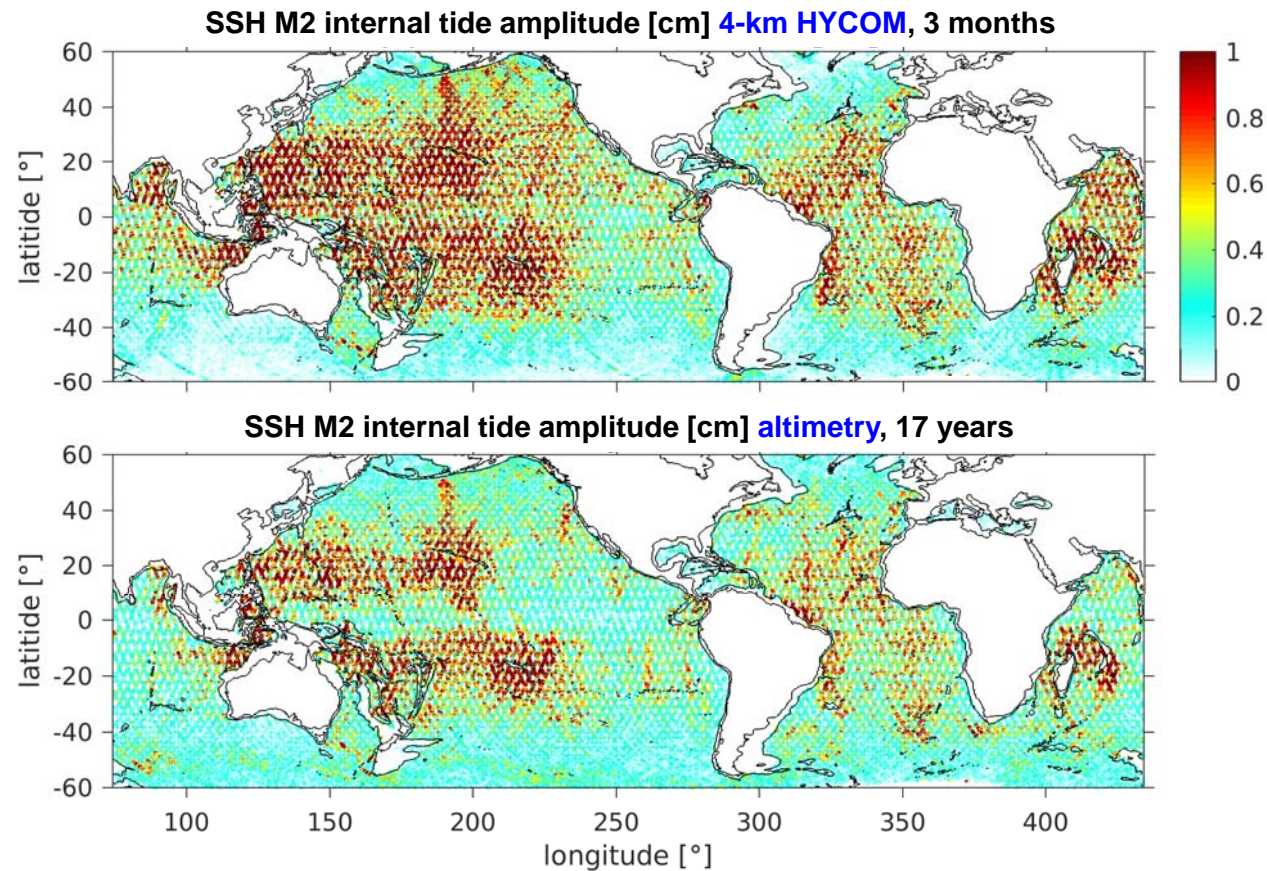


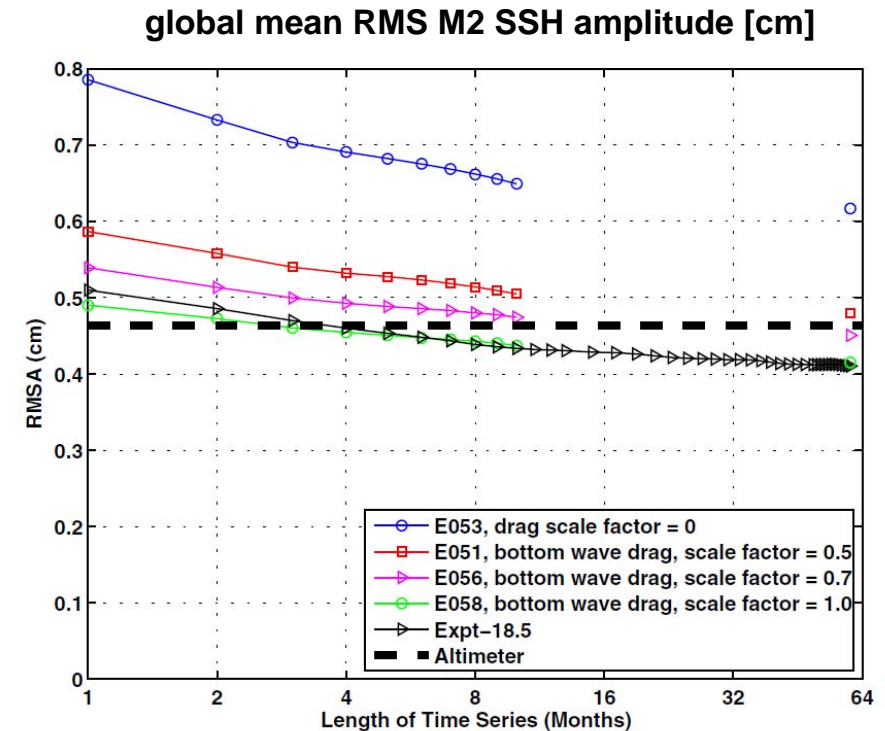
A Correction Factor for Internal Tide Variance in Models



Why is the M_2 internal tide SSH amplitude of a short HYCOM record larger than the amplitude extracted from altimetry?

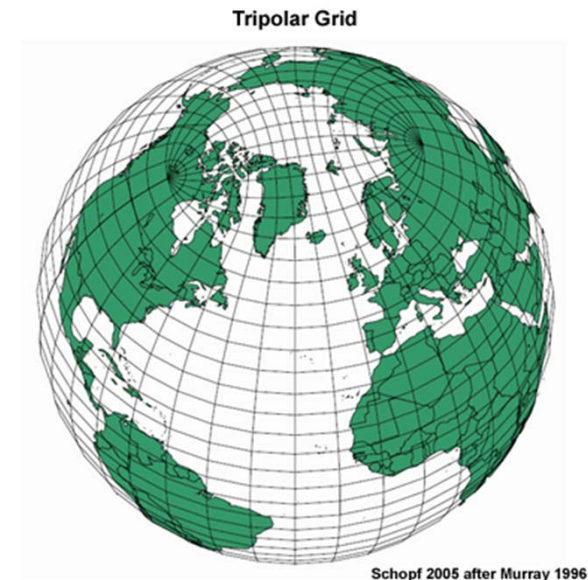
Objectives

- The internal tide M_2 SSH amplitude decreases as a function of time series duration due to refraction/reflection/ducting by mesoscale background flow
- We will compute the SSH variance decay for each horizontal grid cell for a 6-year long HYCOM simulation
- We will use this decay relation to compute a spatially varying correction factor
- The correction factor can be applied to shorter duration simulations to ensure an “apples to apples” comparison to altimetry



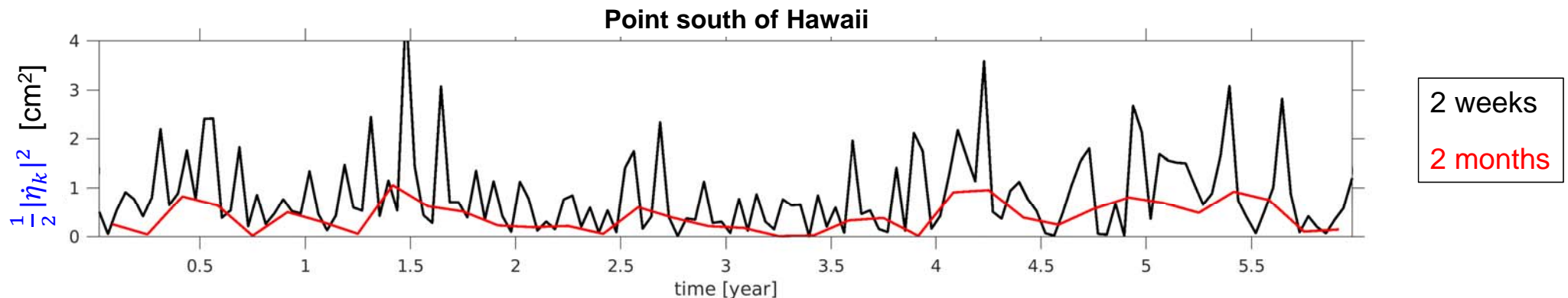
Methodology

- We analyze the **HY**brid **C**oordinate **O**cean **M**odel (HYCOM) simulation Expt_185:
 - Forward simulation
 - 8-km horizontal resolution and 32 layers
 - atmospheric forcing from FNMOC NOGAPS
 - forced with M_2 , S_2 , N_2 , K_2 , K_1 , O_1 , P_1 , Q_1 tidal constituents and SAL correction
 - 6-year simulation; data output every hour
- Steric SSH time series are split into time series with durations of 2 weeks, 1, 2, 3, 4, 6, 9, 12 months, 2, 3, 4, 6 years
⇒ 2-week series has 144 members, 1 month has 72, etc.
- A least squares-harmonic analysis is applied to each member k to extract complex M_2 internal tide amplitude $\dot{\eta}_k = a_k + ib_k$



Total and stationary M₂ tide

- Complex amplitude $\dot{\eta}_k = a_k + ib_k$
- The mean total variance over all members (6 years): $\overline{\frac{1}{2}|\dot{\eta}_k|^2} = \frac{1}{M} \sum_{k=1}^M \frac{1}{2}|\dot{\eta}_k|^2$
- The mean stationary variance over all members: $\frac{1}{2}|\overline{\dot{\eta}_k}|^2$
- Non-stationary fraction: $\frac{\overline{|\dot{\eta}_k|^2} - |\overline{\dot{\eta}_k}|^2}{\overline{|\dot{\eta}_k|^2}}$
- M₂ tidal variance decreases with time series duration:

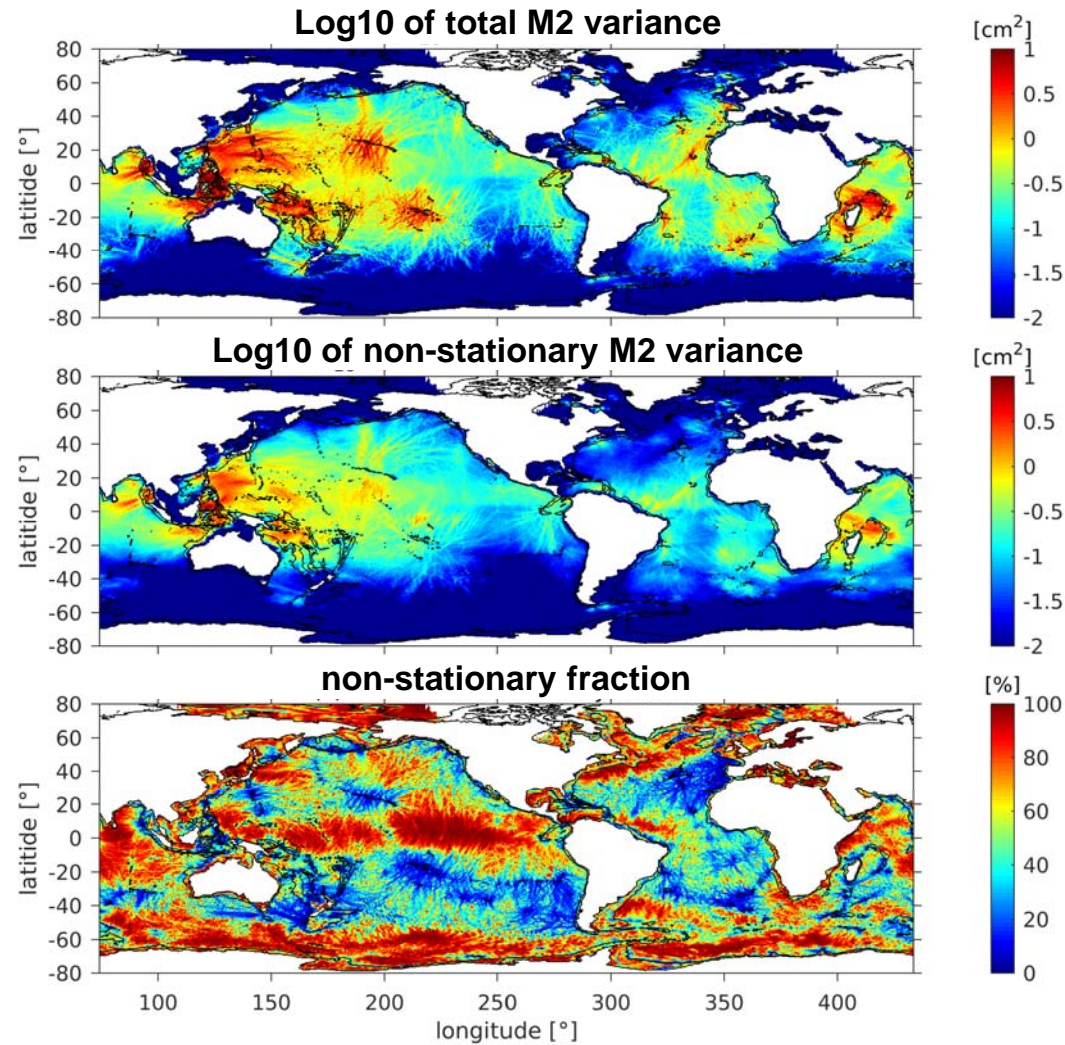


Variance for bi-weekly series

$$\overline{\frac{1}{2} |\dot{\eta}_k|^2}$$

$$\frac{1}{2} (\overline{|\dot{\eta}_k|^2} - \overline{|\dot{\eta}_k|^2})$$

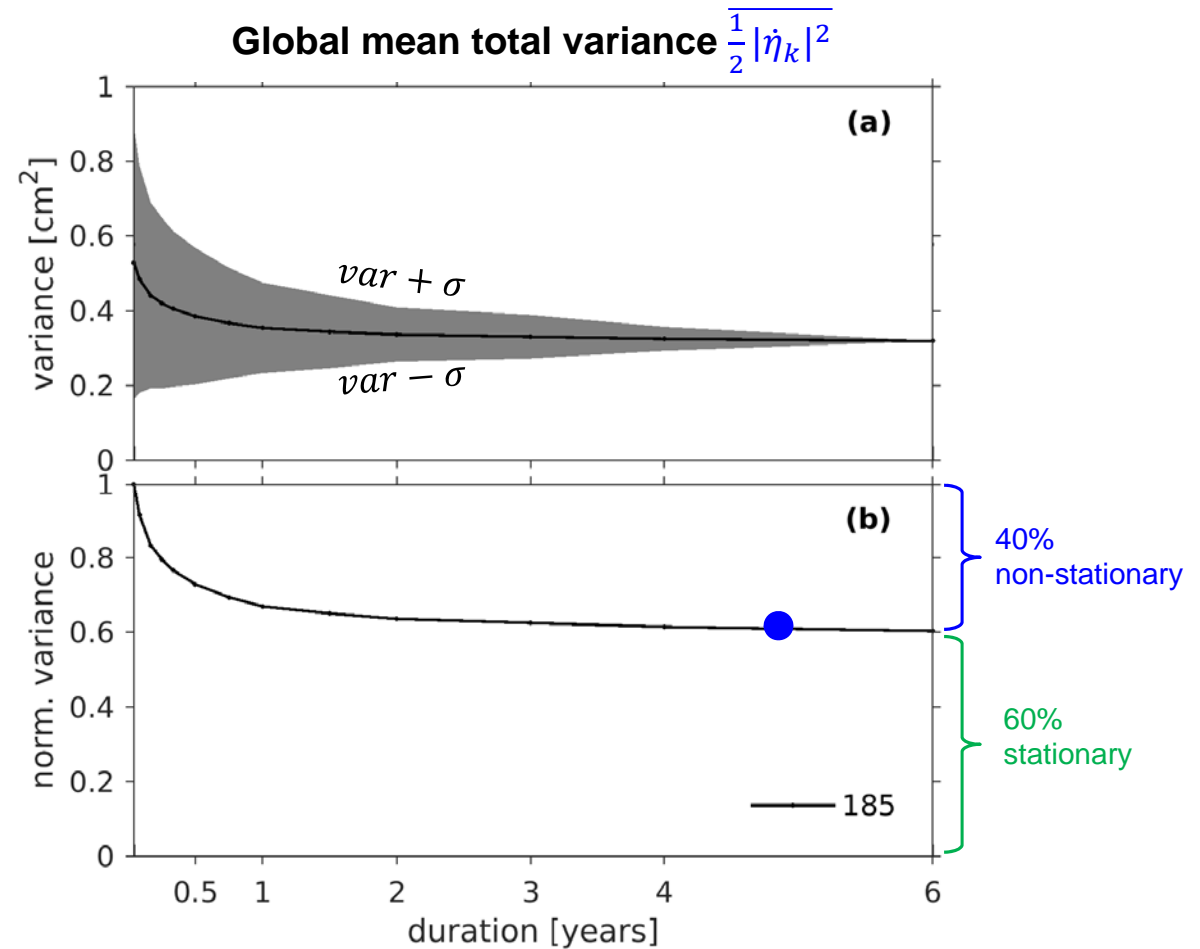
$$\frac{\overline{|\dot{\eta}_k|^2} - \overline{|\dot{\eta}_k|^2}}{\overline{|\dot{\eta}_k|^2}}$$



Variance as a function of duration

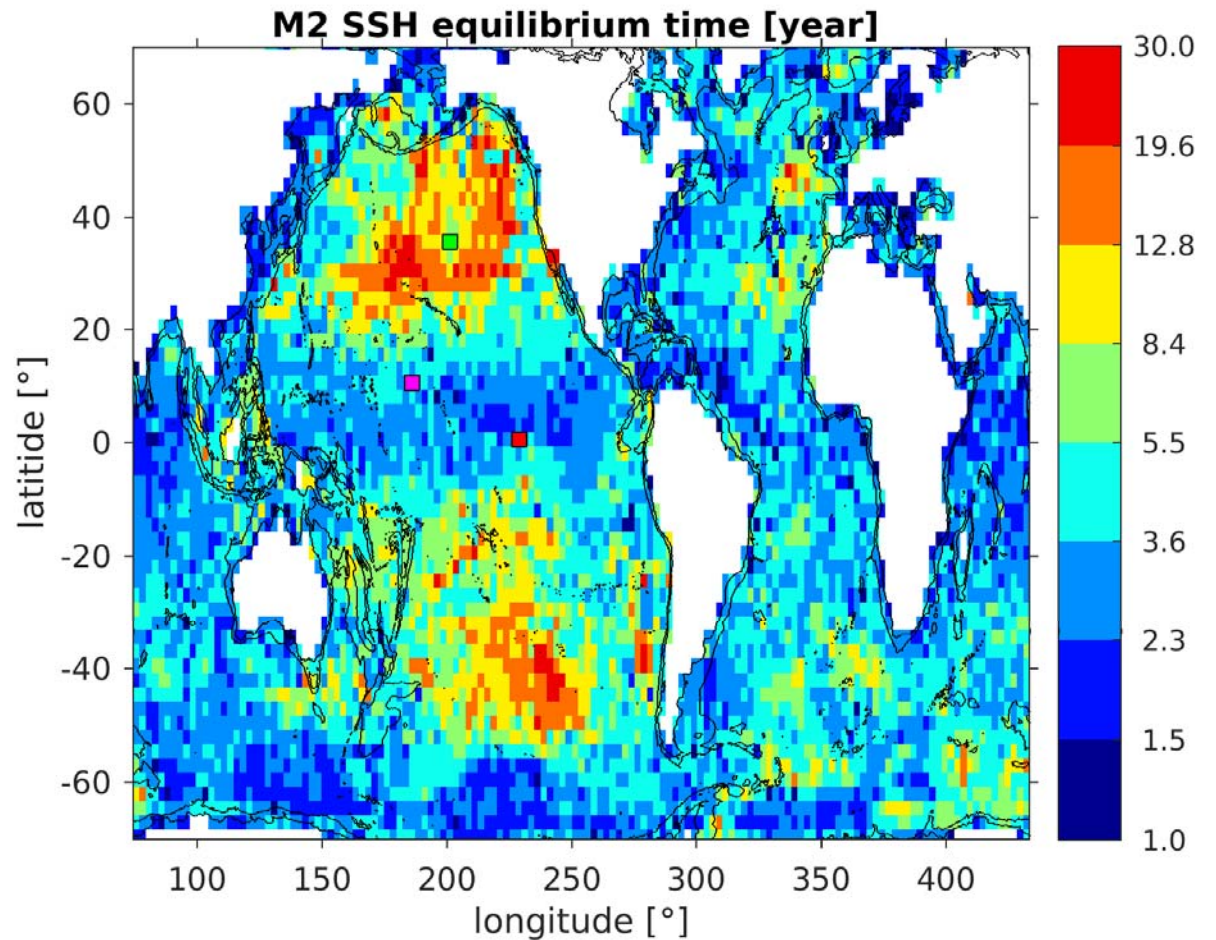
- Total M_2 variance decreases with duration
- **Equilibrium time (●)**
is the time at which the decay rate has dropped below a “threshold” value

 \Rightarrow the variance no longer decays beyond T_{eq}

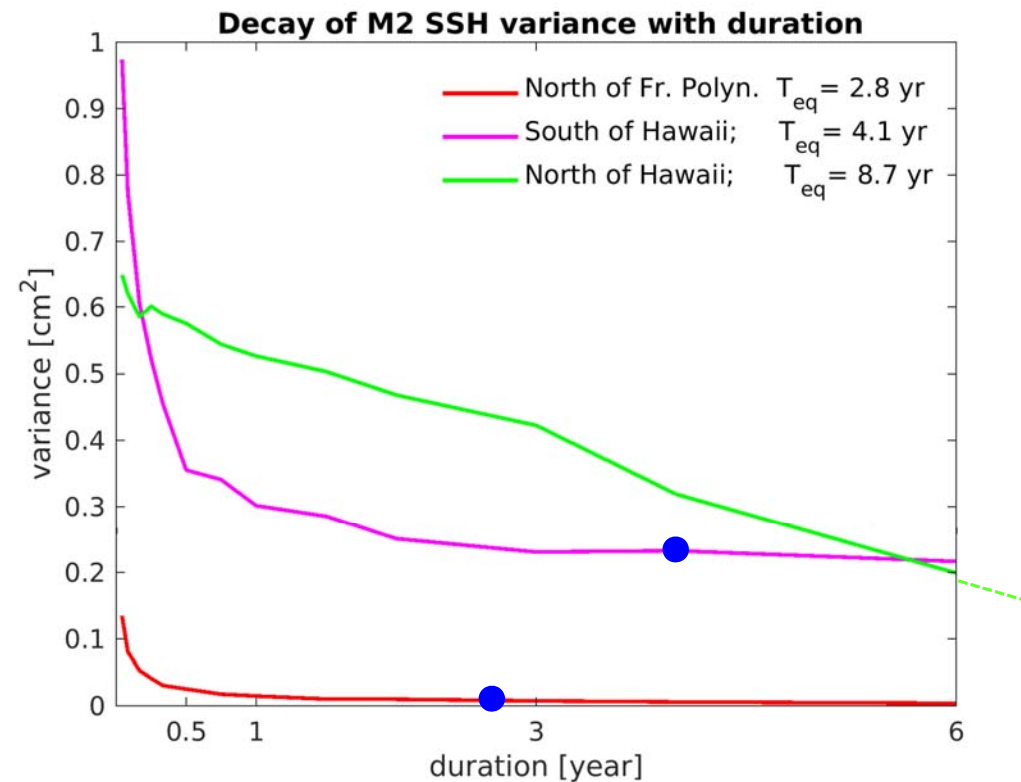
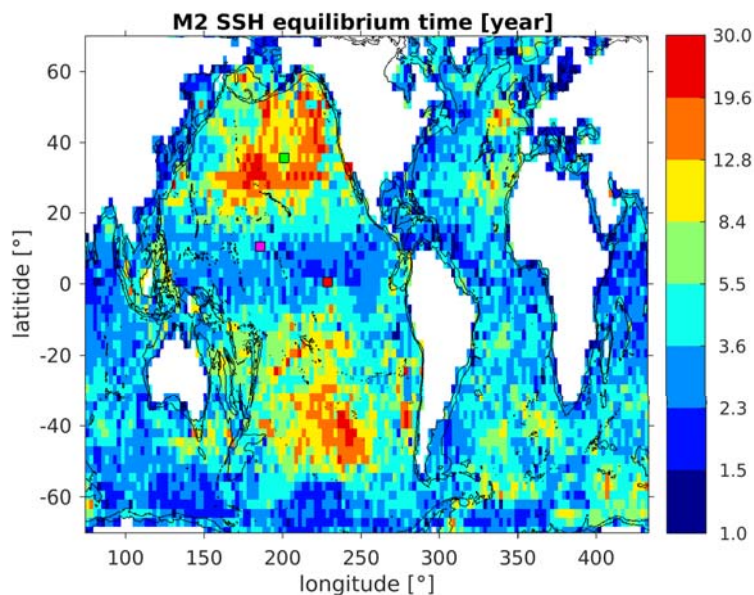


Equilibrium time T_{eq}

- Variance is averaged to 2.5° bins
- $T_{eq} < 6$ years in tropics and Western Boundary currents
- $T_{eq} > 6$ years in subtropical gyres



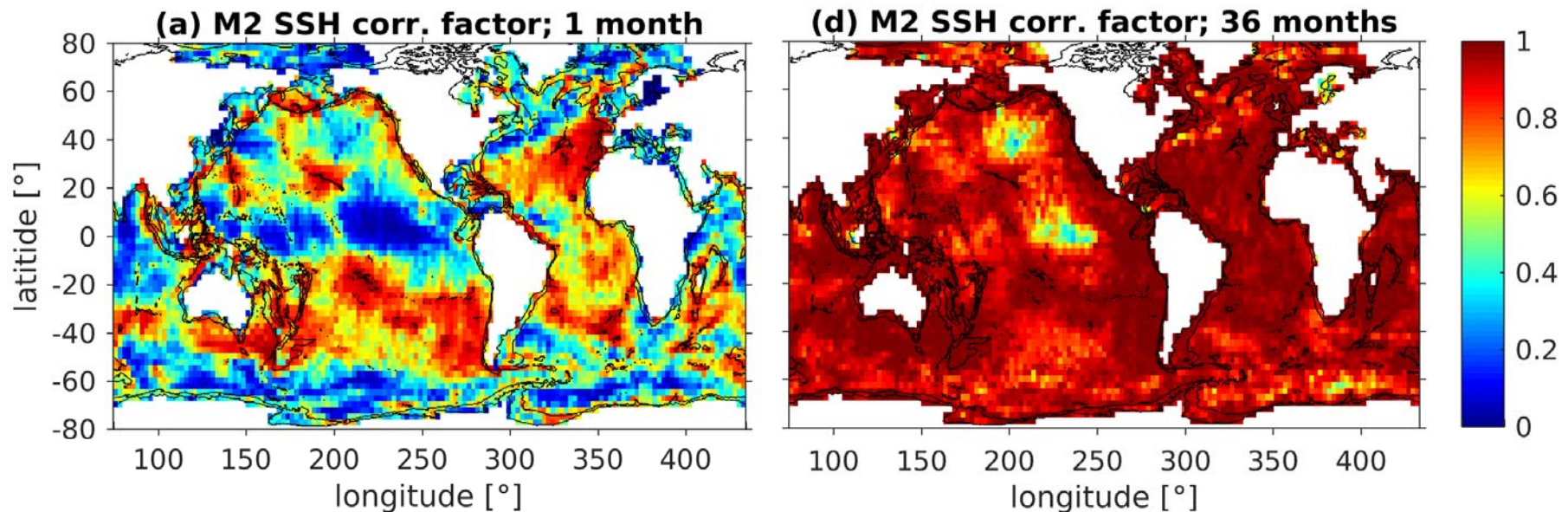
Variance decay in 3 Locations



- Variance in tropics equilibrates quickly
- Gyres do not appear to equilibrate in HYCOM!!

Correction factor $\frac{var_{eq}}{var_{dur}}$

- We can use the decay relations for expt_185 to compute a spatially varying correction factor
- The correction factor is a function of duration
- Relatively short-duration simulations can be corrected in order to compare them with altimetry

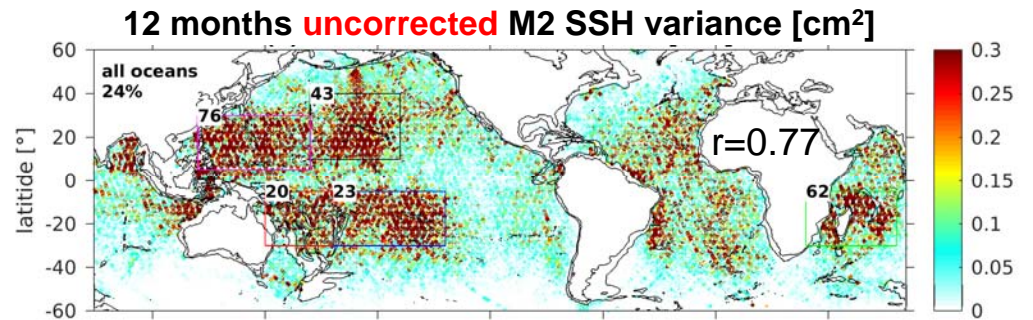
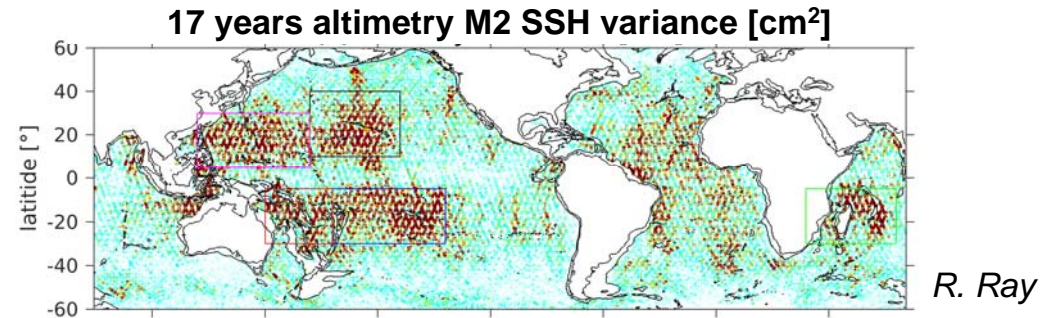
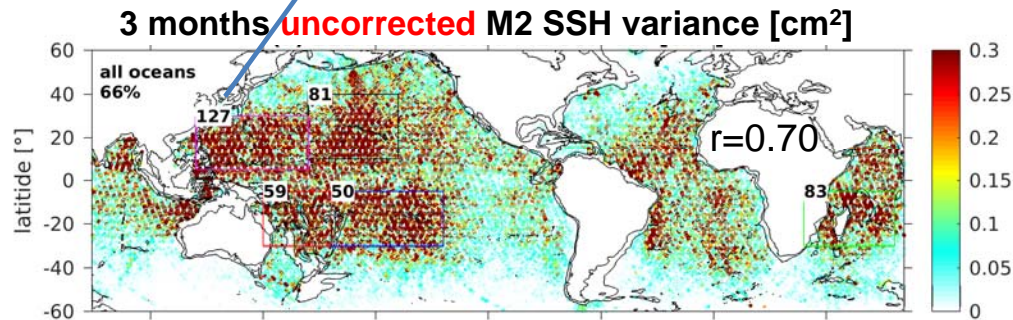


Correction of expt_221

- Correction is applied to 4-km HYCOM expt_221

- Harmonic analysis based on 3 months
- Harmonic analysis based on 12 months

numbers are % difference with altimetry



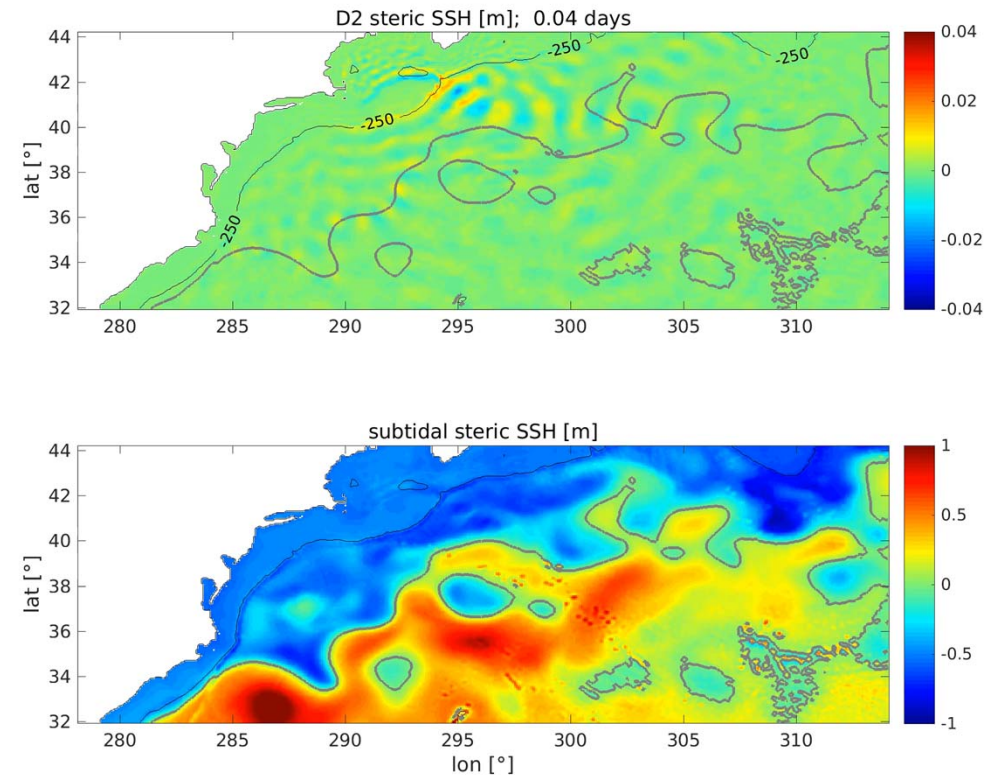
Conclusions

- Internal tides become non-stationary in the tropics and western boundary currents
- The tidal variance decay and equilibrium times are spatially variable
- The variance decay can be used to correct SSH variance of shorter simulations

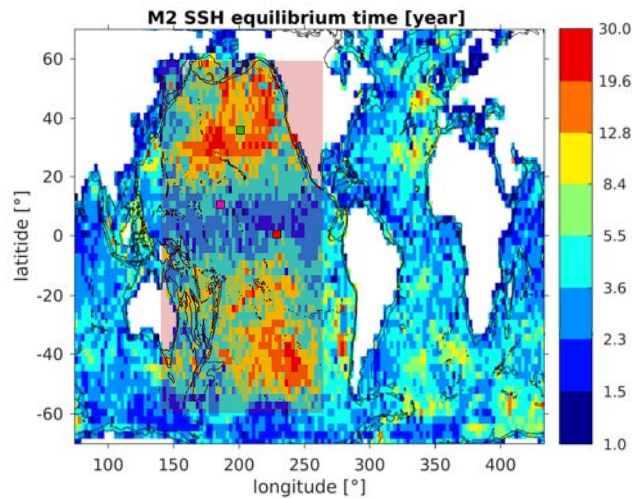
Future work:

- Update the altimetry data for this analysis
- Perform a complex modulation on the altimetry data to verify variance decay in the gyres
- Extend the 1-year 4-km simulation to 6 years

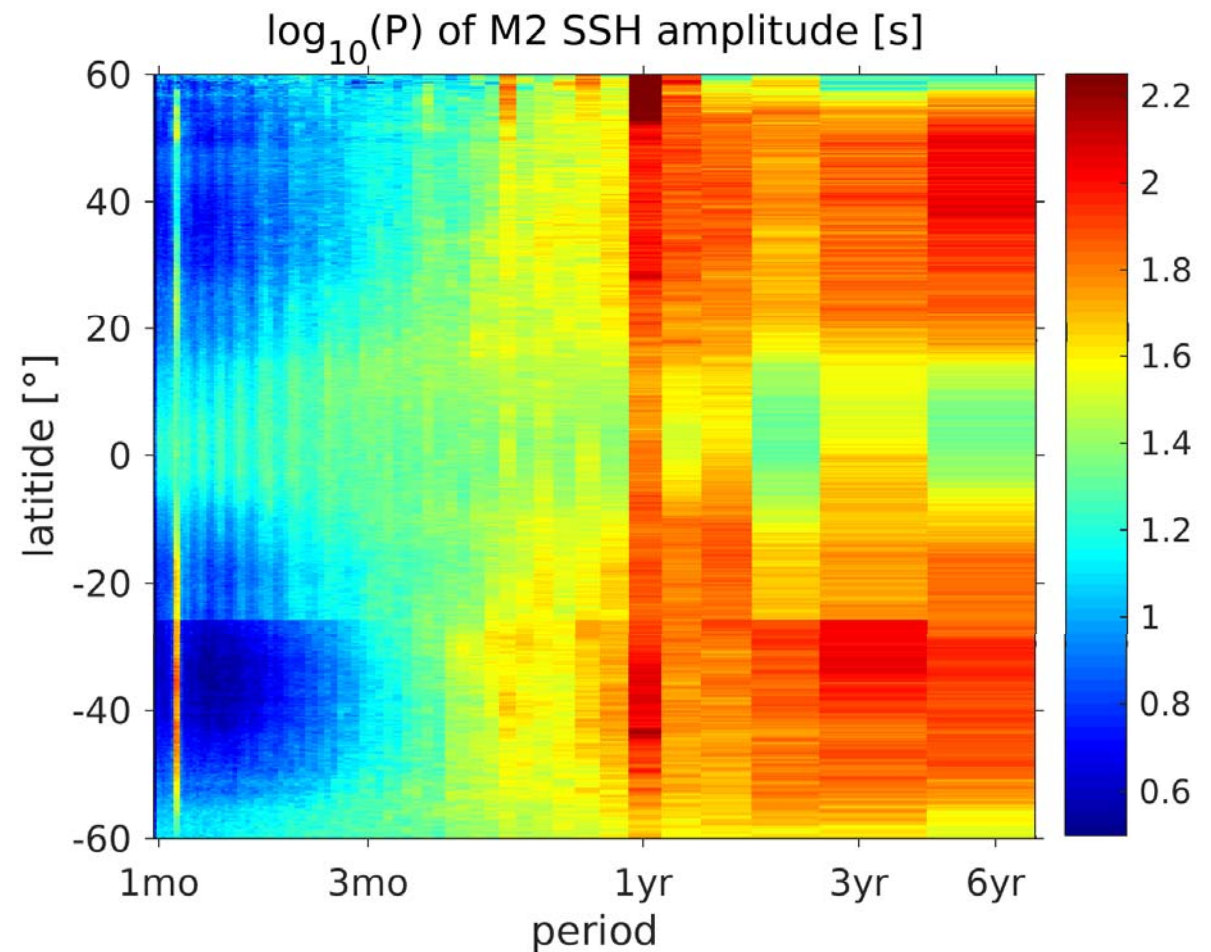
M_2 internal tides from Georges Bank



Why no equilibrium in the gyres?

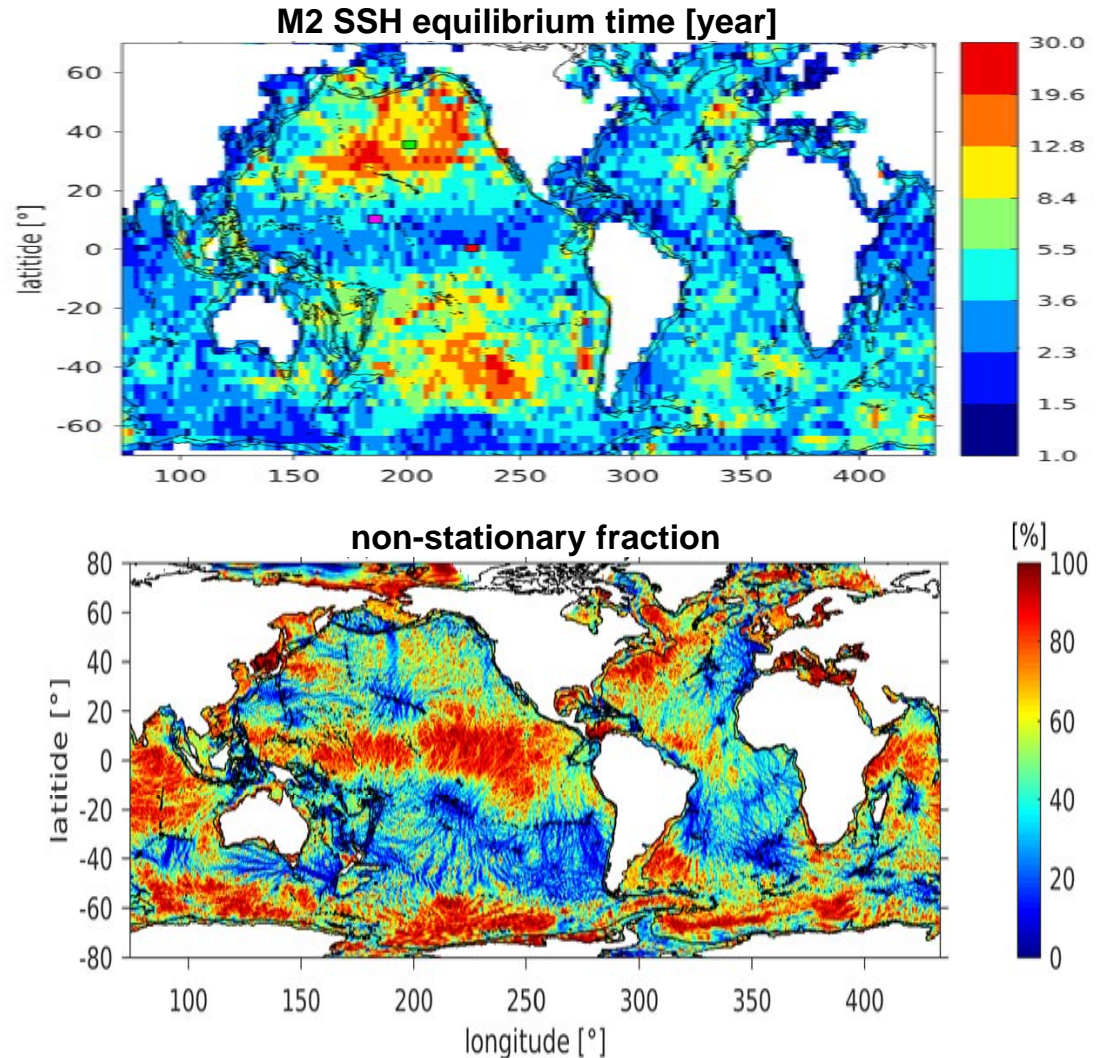


- Compute FFT for the normalized amplitude of the bi-weekly time series with 144 members
- In gyres, modulation periods longer than 1 yr have much energy

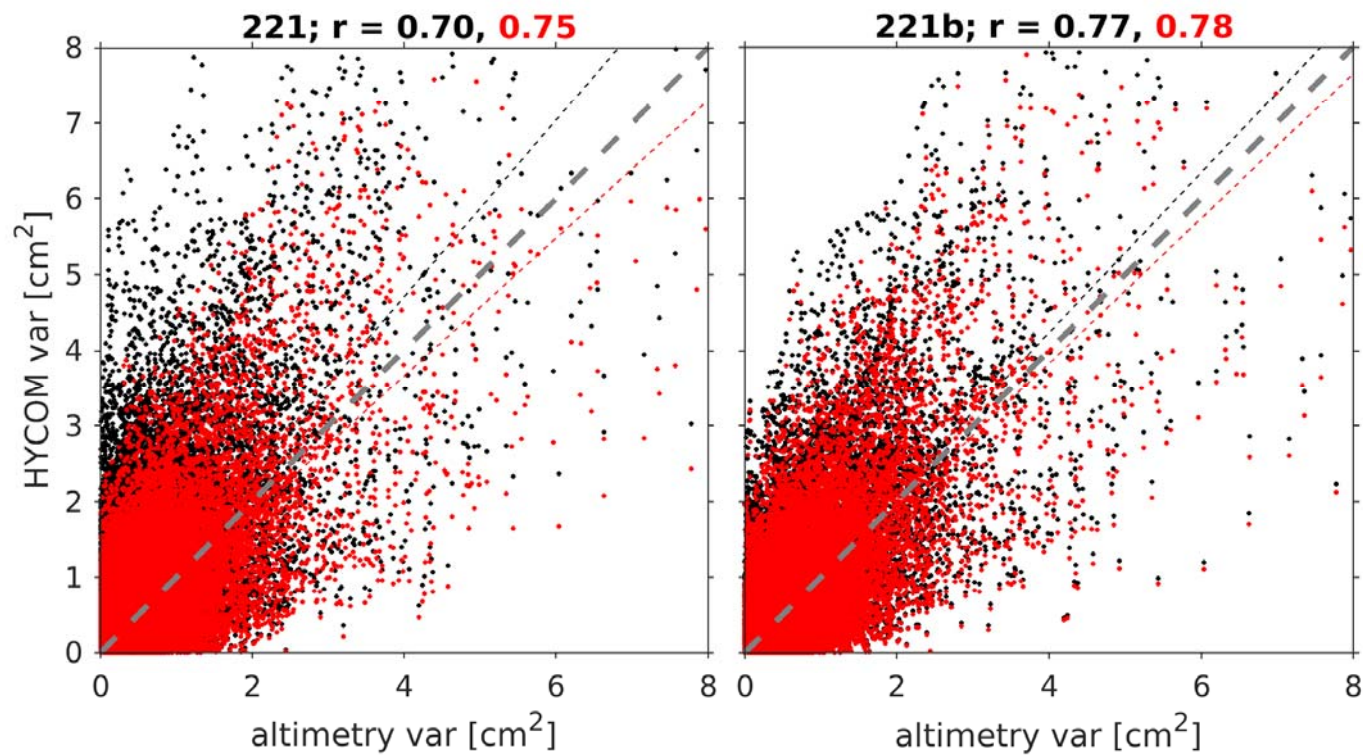


Equilibrium time T_{eq}

- Variance is averaged to 2.5° bins
- $T_{eq} < 6$ years in tropics and Western Boundary currents
- $T_{eq} > 6$ years in subtropical gyres
- The smaller T_{eq} , the larger the non-stationary fraction

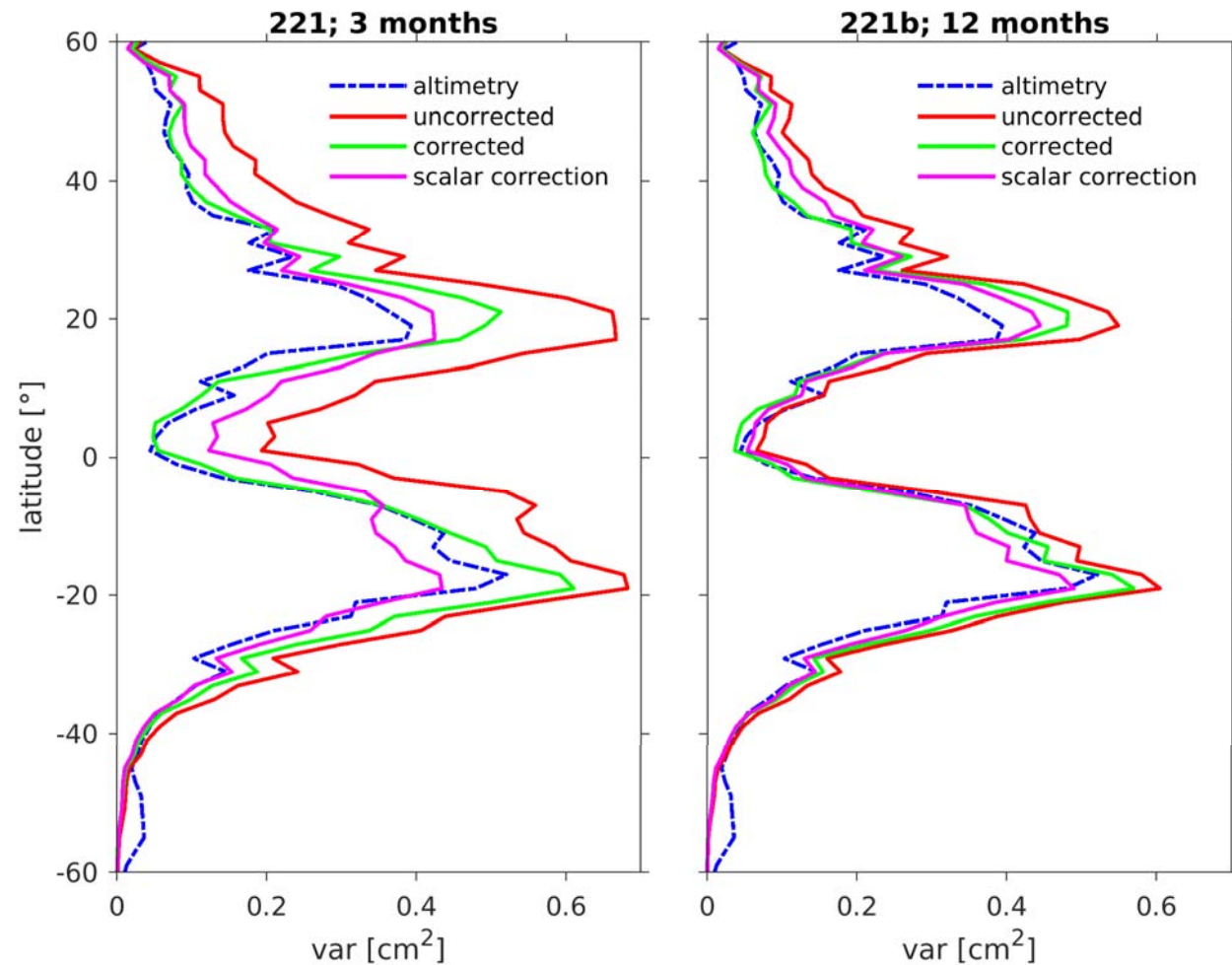


Scatter plot



Meridional Transect Across Pacific

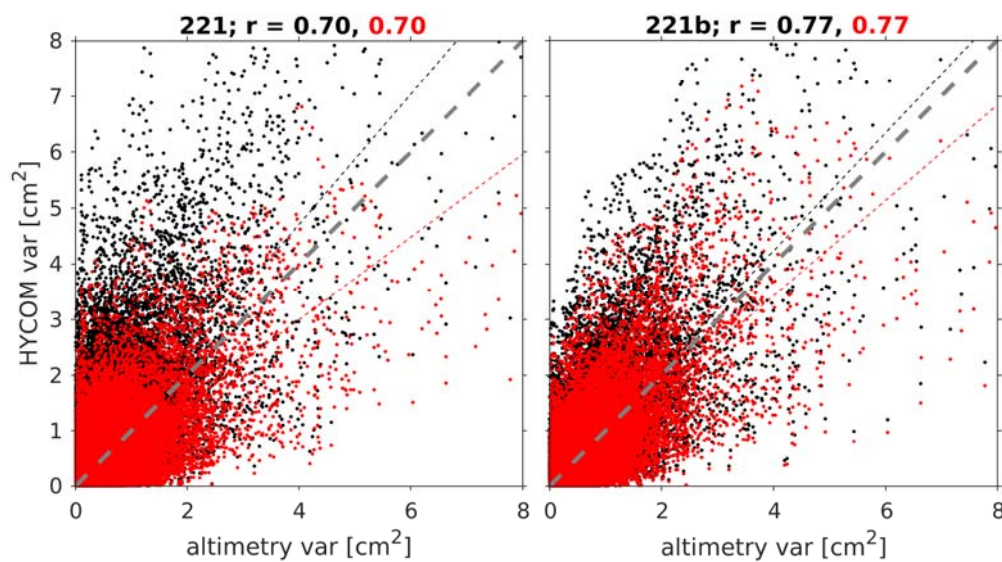
- Global-mean scalar correction
 - 3 months: 0.63
 - 12 months: 0.8
- Scalar correction
 - reduces variance closer to altimetry
 - But yields a weaker spatial correlation



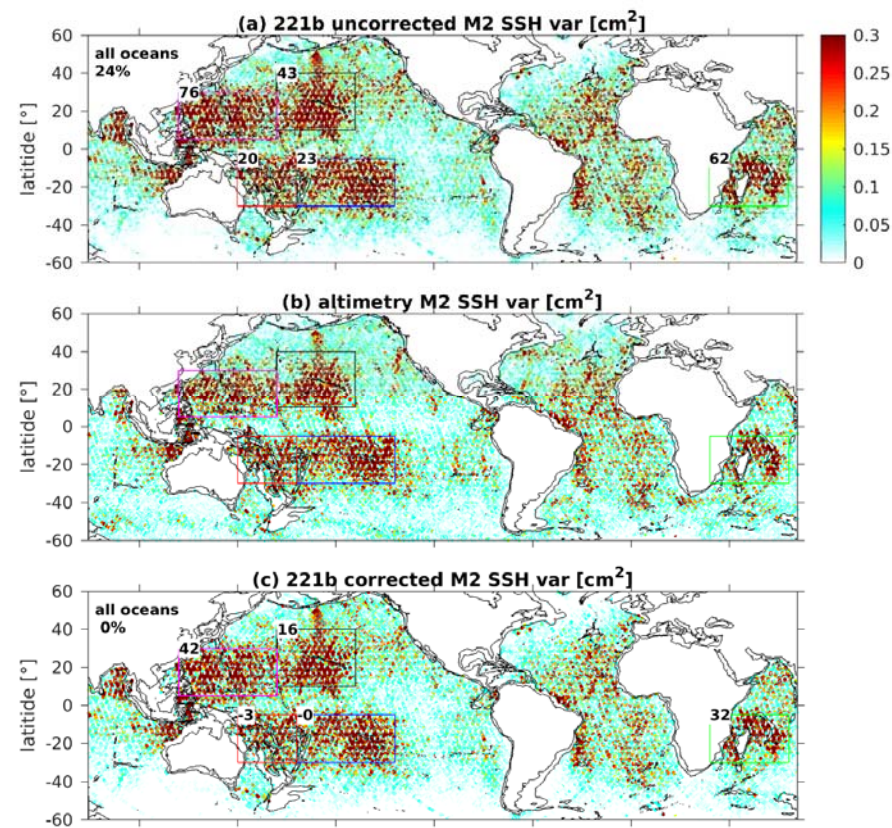
Scalar correction

3 months

12 months



12 months



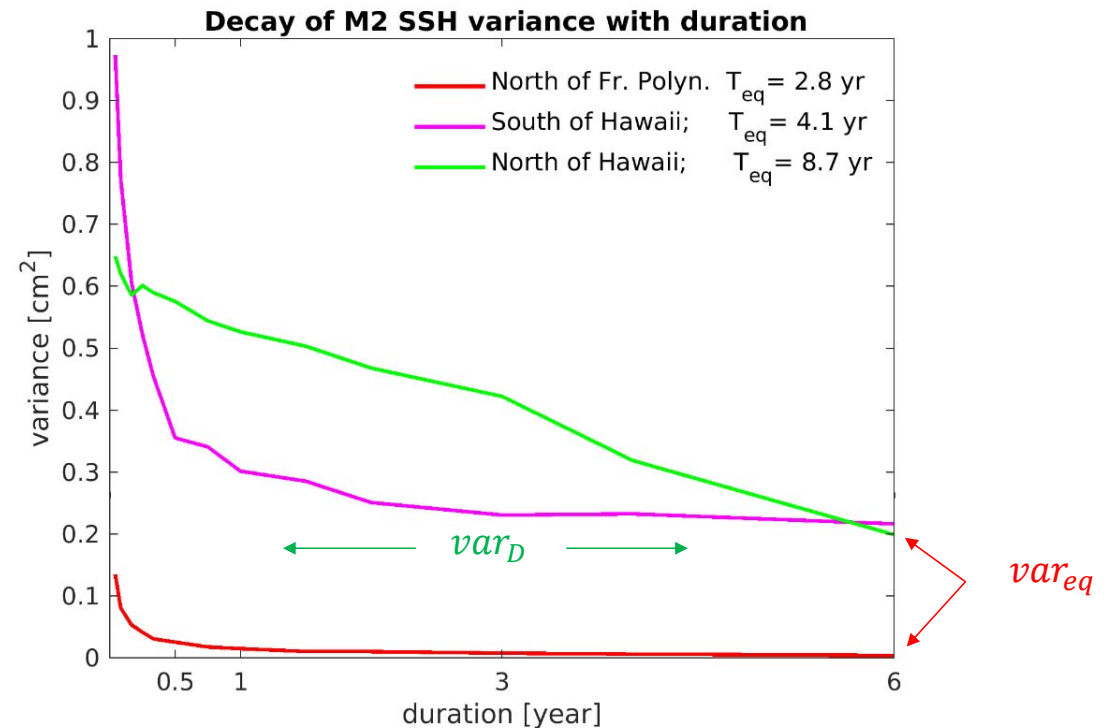
Compute Correction

- How can we compare HYCOM with satellite altimetry?
- We assume
 - equilibrium variance values are reached by 6 years in HYCOM
 - decay process in HYCOM is realistic

- Equilibrium variance of a data set:

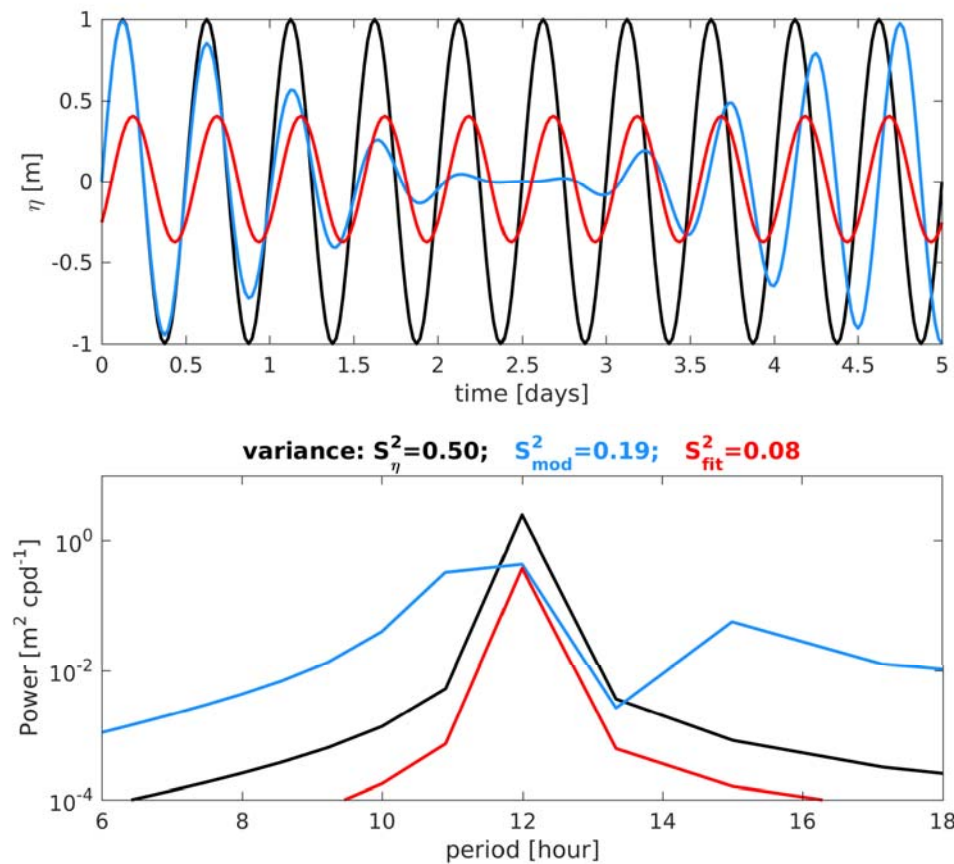
$$var_{eq,data} = \frac{var_{eq}}{var_D} var_{D,data}$$

- Variance of data set: $var_{D,data}$ with duration D
 - Variance of HYCOM for duration D : var_D
 - Equilibrium variance in HYCOM: var_{eq}
- $var_{eq,data}$ can be compared with altimetry variance



Non-stationarity affects variance

- Internal tides are modulated by the time variability of the background flow
- The stronger the modulation, the smaller the variance of **the least-squares fit**



Show “decay rate”