

# CNES/GRGS gravity field solutions from GRACE: RL03-v2

*J.-M. Lemoine <sup>(1)</sup>, S. Bourgogne <sup>(2)</sup>, S. Bruinsma <sup>(1)</sup>, F. Reinquin <sup>(1)</sup>, P. Gégout <sup>(3)</sup>, R. Biancale <sup>(1)</sup>*

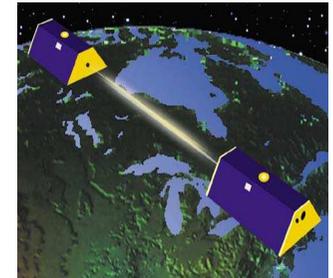
*(1) CNES/GRGS, Toulouse, France*

*(2) Géode & Cie, Toulouse, France*

*(3) GET/UMR5563/OMP/GRGS, Toulouse, France*

## GRACE (L-1B “Version2” data)

- K-Band Range-Rate data ( $\sigma_{\text{apriori}} = .1 \mu\text{m}$ )
- Accelerometer / attitude
- GPS data (1-day arcs,  $\sigma_{\text{code}} = .8 \text{ m}$ ,  $\sigma_{\text{phase}} = 20 \text{ mm} / 30\text{s resolution}$ )  
(actually:  $\sigma_{2002-2003} = 8 \text{ mm}/30 \text{ s}$ ,  $\sigma_{2003-2013} = 20 \text{ mm}/300 \text{ s}$ ,  $\sigma_{2013-2015} = 8 \text{ mm}/30 \text{ s}$ )



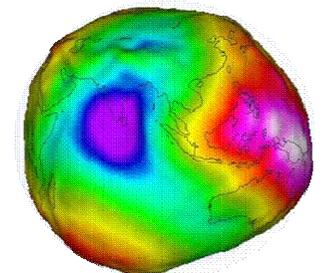
## SLR

- Lageos1/2 data (10-day arcs,  $\sigma_{\text{apriori}} = 6 \text{ mm}$ )
- Starlette/Stella data (5-day arcs,  $\sigma_{\text{apriori}} = 10 \text{ mm}$ )



## Physical parameters present in the normal equations

- Gravity spherical harmonic coefficients complete to **degree and order 175** (truncated to **30 for LAGEOS** and **40 for GPS** data)
- Ocean tides s. h. coefficients for 14 tidal waves with maximum degree/order  $\leq 30$  (not used yet)



# Models used : v0 → v2



## Dynamical models

Gravity	<i>EIGEN-GRGS.RL02</i> → <i>EIGEN-6S2 (LAGEOS/GRACE/GOCE)</i>
Ocean tide	<i>FES2004 (degree 80)</i> → <i>FES2012 (Legos)</i>
Atmosphere	<i>3-D ECMWF pressure grids / 6hrs</i> → <i>ERA-interim / 3hrs</i>
Ocean mass model	<i>MOG2D (non-IB) / 6hrs</i> → <i>TUGO (Legos) / 3hrs</i>
Atmospheric tides	→ <i>Not necessary any more</i>
3 <sup>rd</sup> body	<i>Sun, Moon, 6 planets (DE405)</i>
Solid Earth tides	<i>IERS Conventions 2010</i>
Pole tides	<i>IERS Conventions 2010</i>
Non gravitational	<i>Accelerometer data (+biases and scale factors)</i>

## Geometrical models

SLR stations	<i>ITRF2008 coordinates</i> → <i>updated</i>
GPS	<i>IGS orbits and CODE clocks</i> → <i>IGS Repro-1 orbits and clocks</i>

## Other models

Hydrology	Taken into account by the a priori gravity field
Glacial Isostatic Adjustment	

## ❖ Inversion technique used for RL03 : truncated Singular Value Decomposition (SVD)

- It is more efficient to solve well chosen linear combinations of coefficients (by truncated SVD) than to solve indistinctly the coefficients (by Cholesky decomposition).
- Demonstration with a normal matrix up to degree/order 80 :
  - 1) Solving for the first 2601 components of the canonical basis (i.e. spherical harmonic coefficients up to degree/order 50)
  - 2) Solving for the first 2601 components of the basis made by the eigenvectors of the normal matrix

# 1) Cholesky decomposition

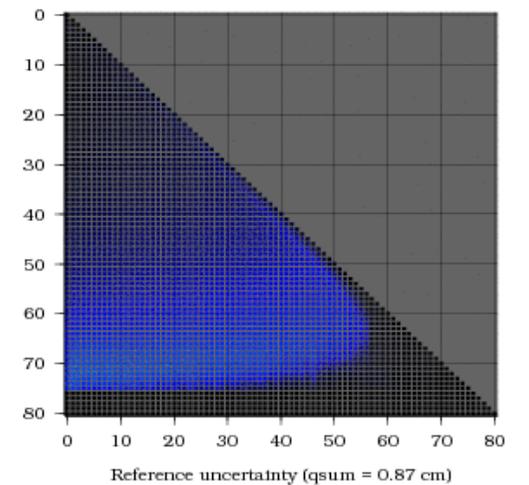
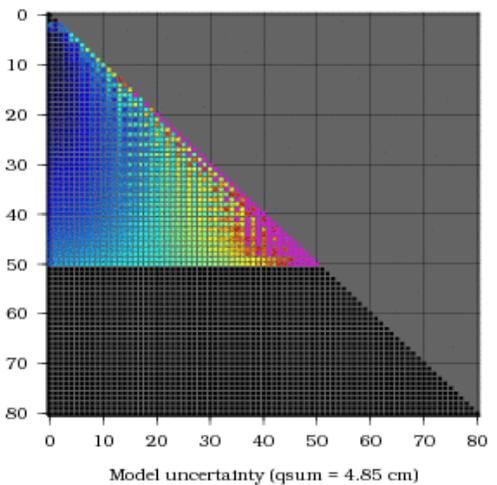
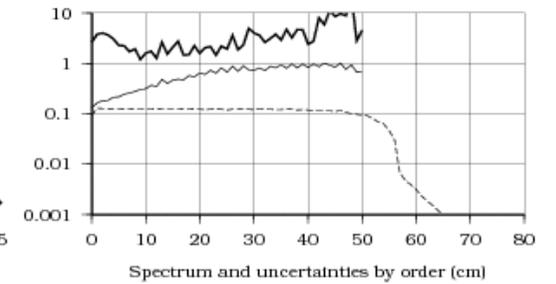
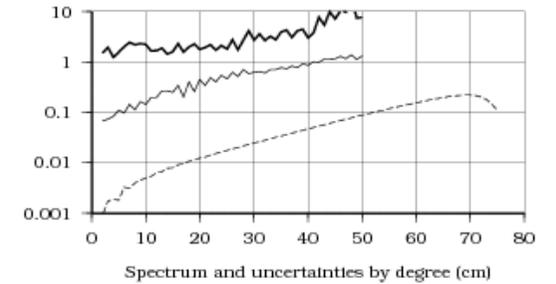
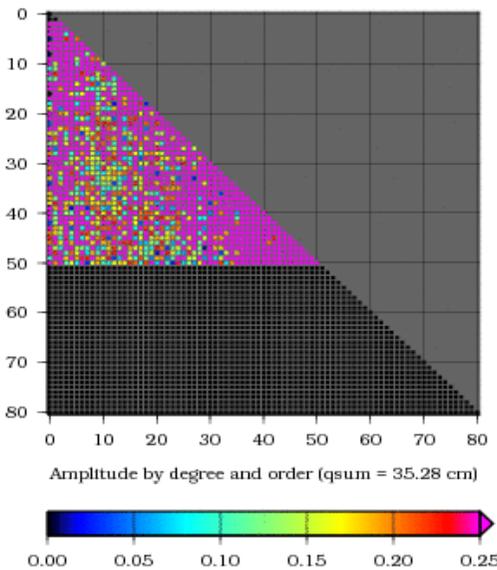
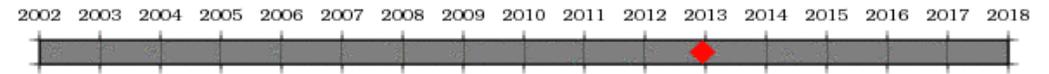
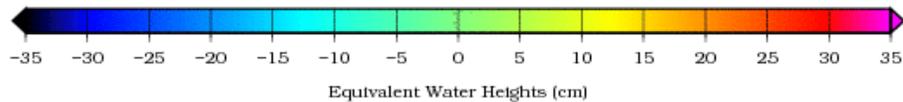
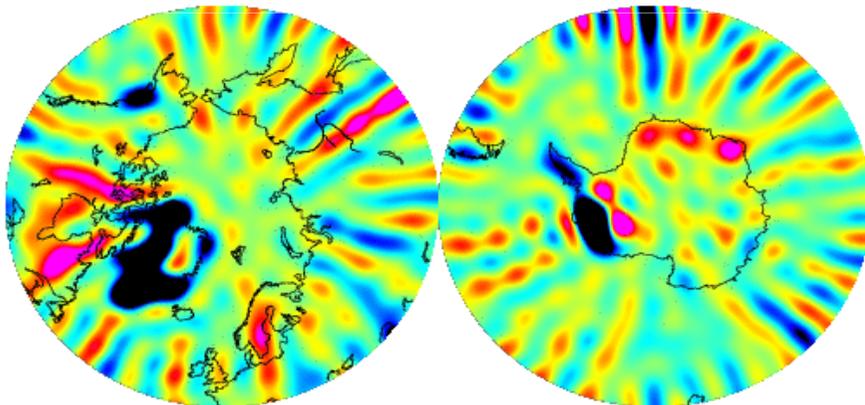
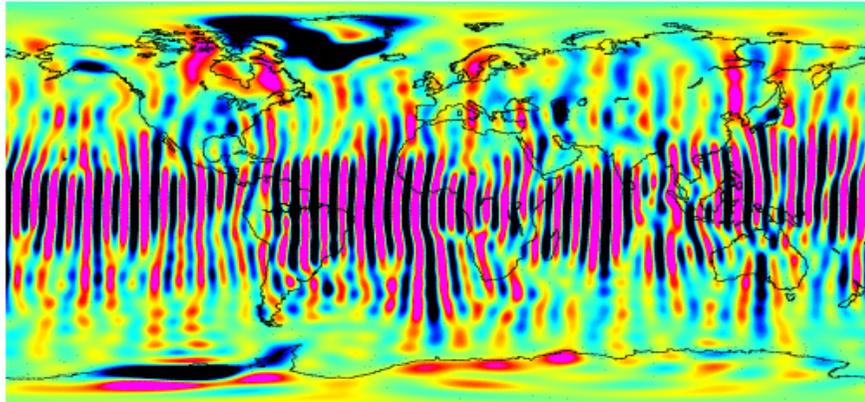
Equivalent Water Heights comparison

Cholesky inversion up to degree and order 50: 2601 parameters

Reference: Mean field

Degree 2 to 80

min -184.81 cm / max 168.34 cm / weighted rms 34.56 cm / oceans 37.61 cm



# 2) Truncated SVD

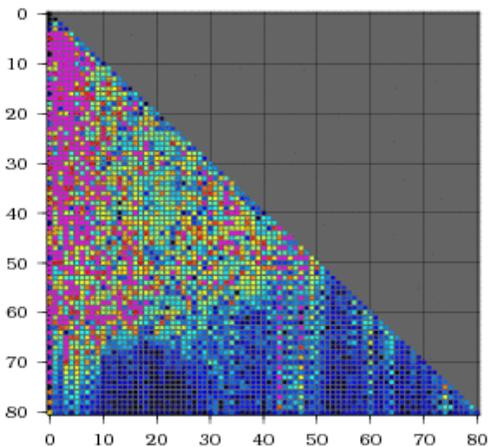
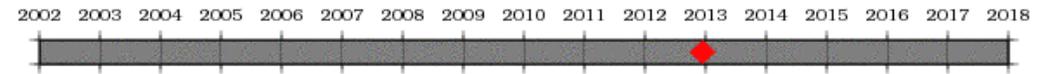
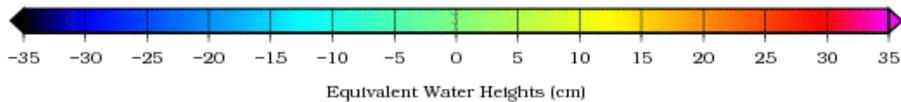
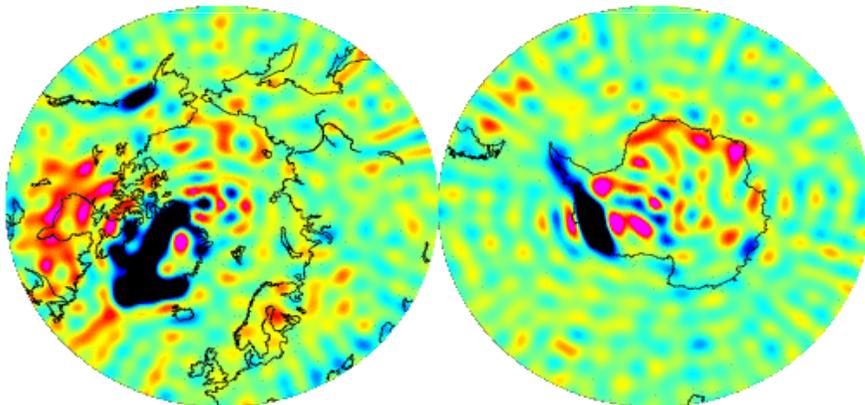
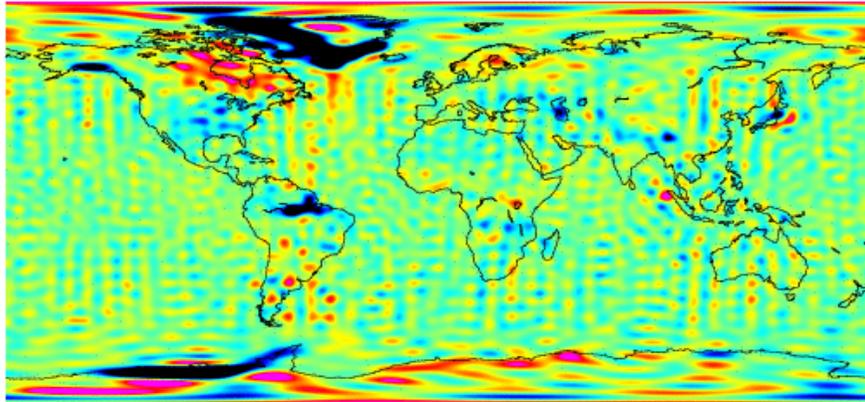
Equivalent Water Heights comparison

SVD solution: minimisation in the direction of the 2601 most significant eigenvectors

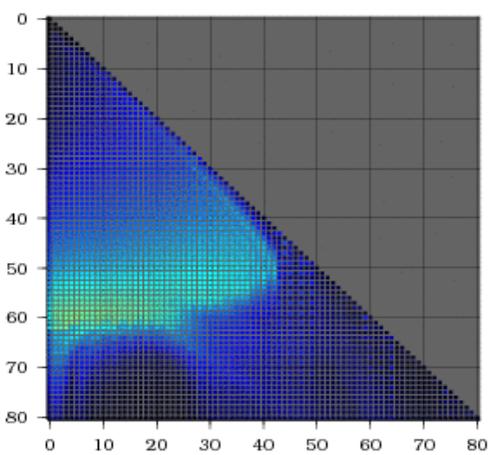
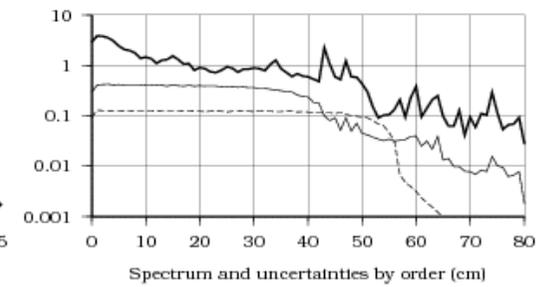
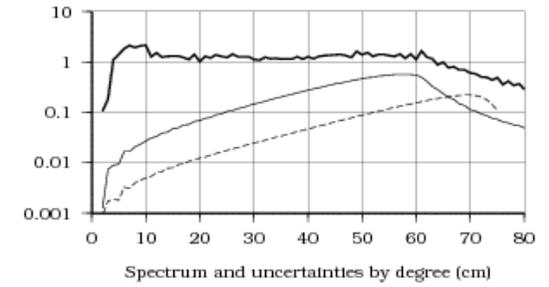
Reference: Mean field

Degree 2 to 80

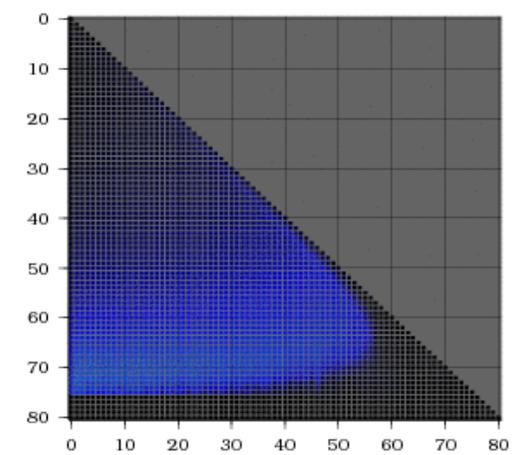
min -206.01 cm / max 58.90 cm / weighted rms 10.72 cm / oceans 6.60 cm



Amplitude by degree and order (qsum = 10.90 cm)  
Spherical Harmonics (cm)



Model uncertainty (qsum = 2.41 cm)



Reference uncertainty (qsum = 0.87 cm)

## ❖ Trying to solve the problems at the poles

- Since SVD does not solve sectorial coefficients due to a lack of information, we need to introduce decent a-priori sectorial coefficients before using SVD
- So we tried to establish a 2-step inversion in RL03-v2
  - First step: Cholesky inversion with constraints to obtain good sectorial coefficients
  - Second step: Truncated SVD inversion starting with the first step solution

## ❖ Results

- The 2-step inversion improves the solutions mainly at the poles

# RL03-v1

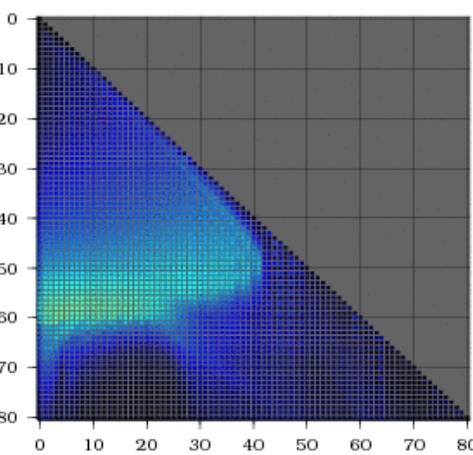
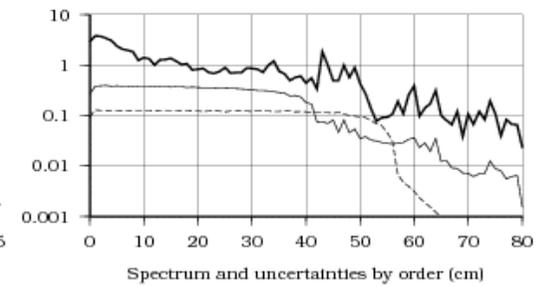
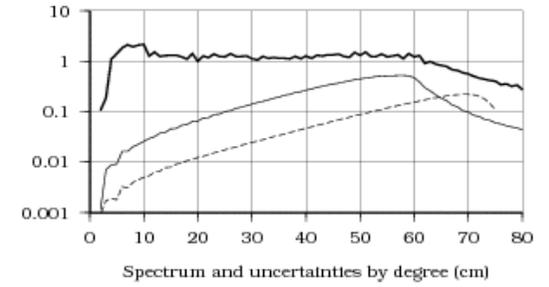
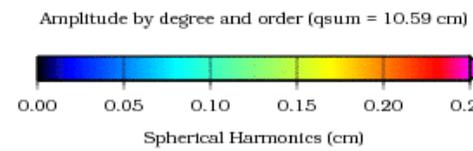
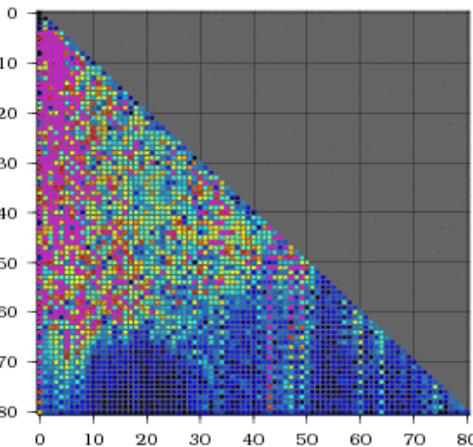
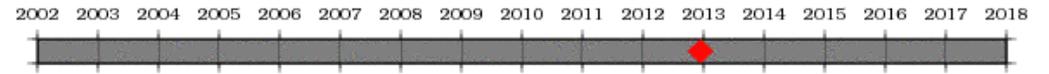
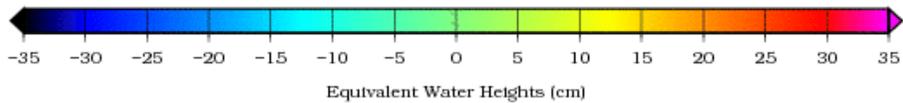
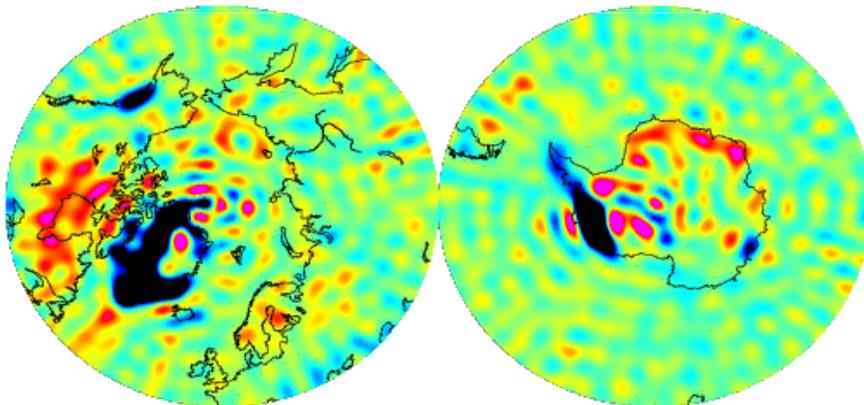
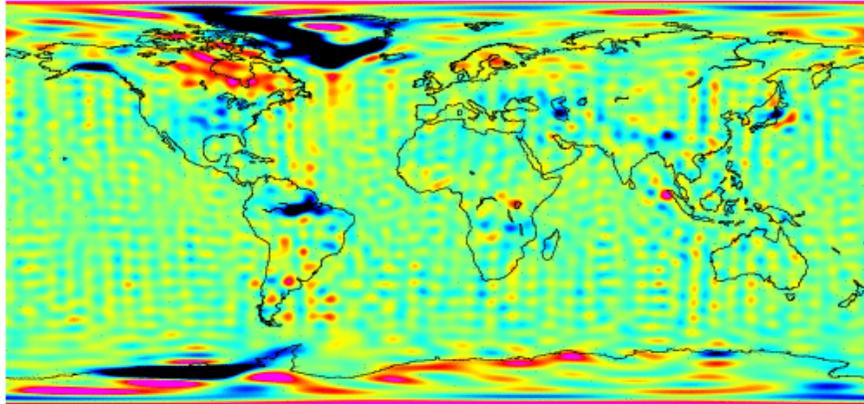
Equivalent Water Heights comparison

T36.decade.22992.O.G\_ONLY.VI\_RL03EQ.svd2500.shc

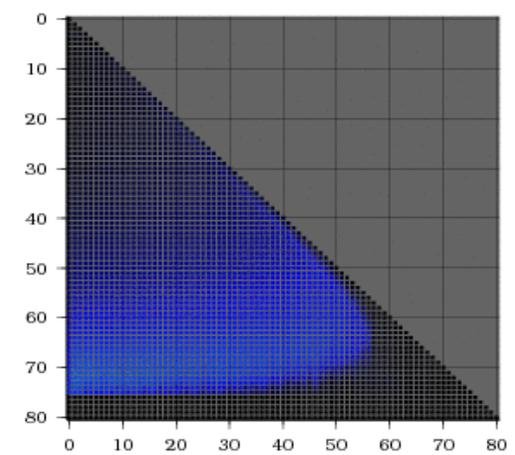
Reference: CHAMP\_MOYEN\_RL03.par\_cumul\_EQN.v2

Degree 2 to 80

min -198.94 cm / max 62.61 cm / weighted rms 10.41 cm / oceans 6.21 cm



Model uncertainty (qsum = 2.22 cm)



Reference uncertainty (qsum = 0.87 cm)

# RL03-v2

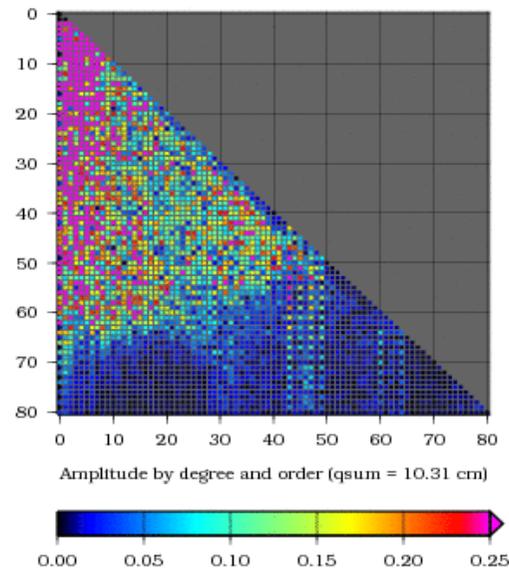
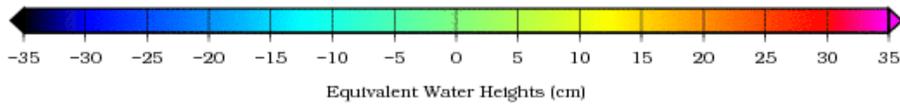
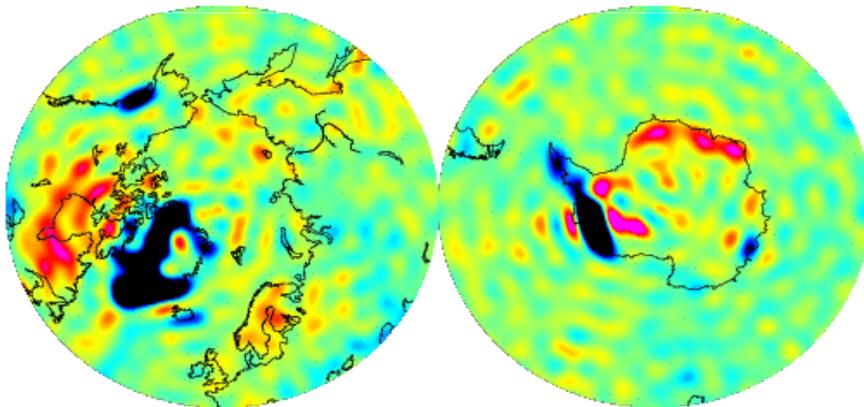
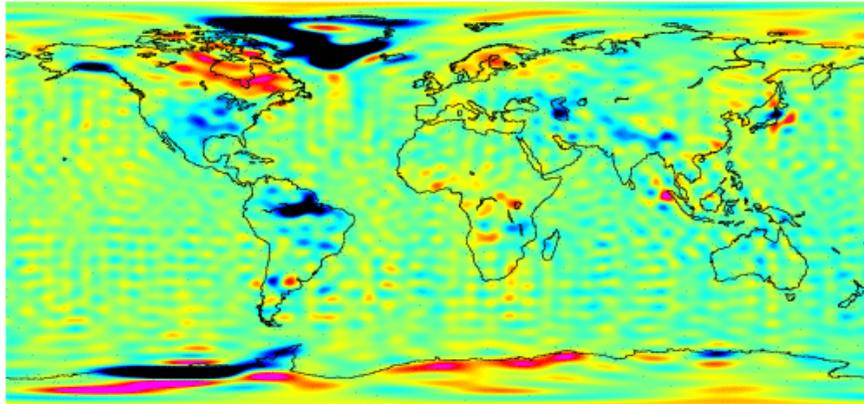
Equivalent Water Heights comparison

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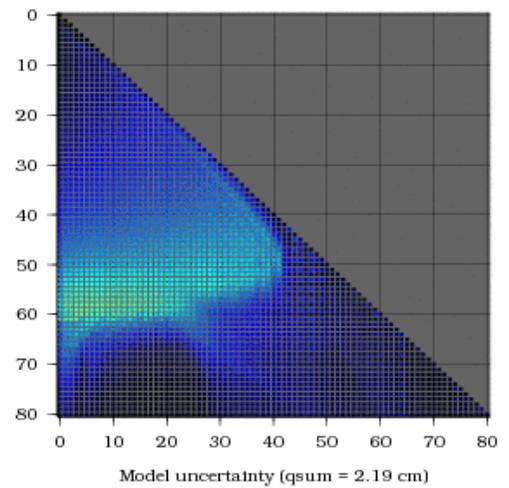
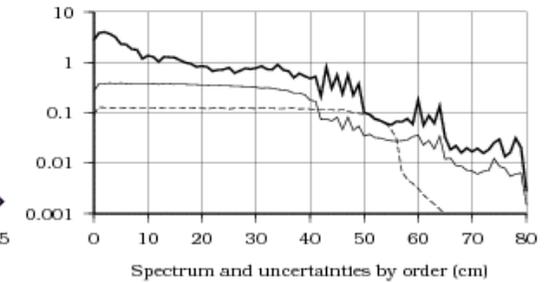
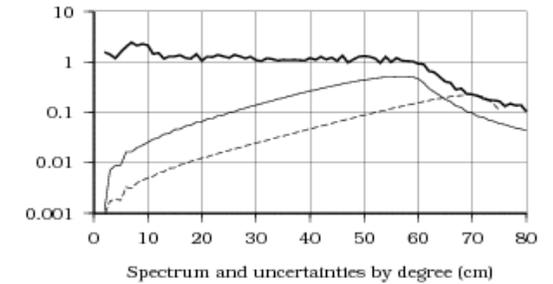
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Degree 2 to 80

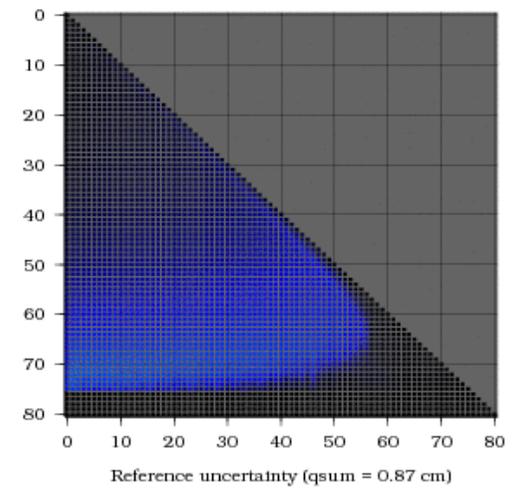
min -206.60 cm / max 55.46 cm / weighted rms 10.18 cm / oceans 5.66 cm



Spherical Harmonics (cm)



Model uncertainty (qsum = 2.19 cm)

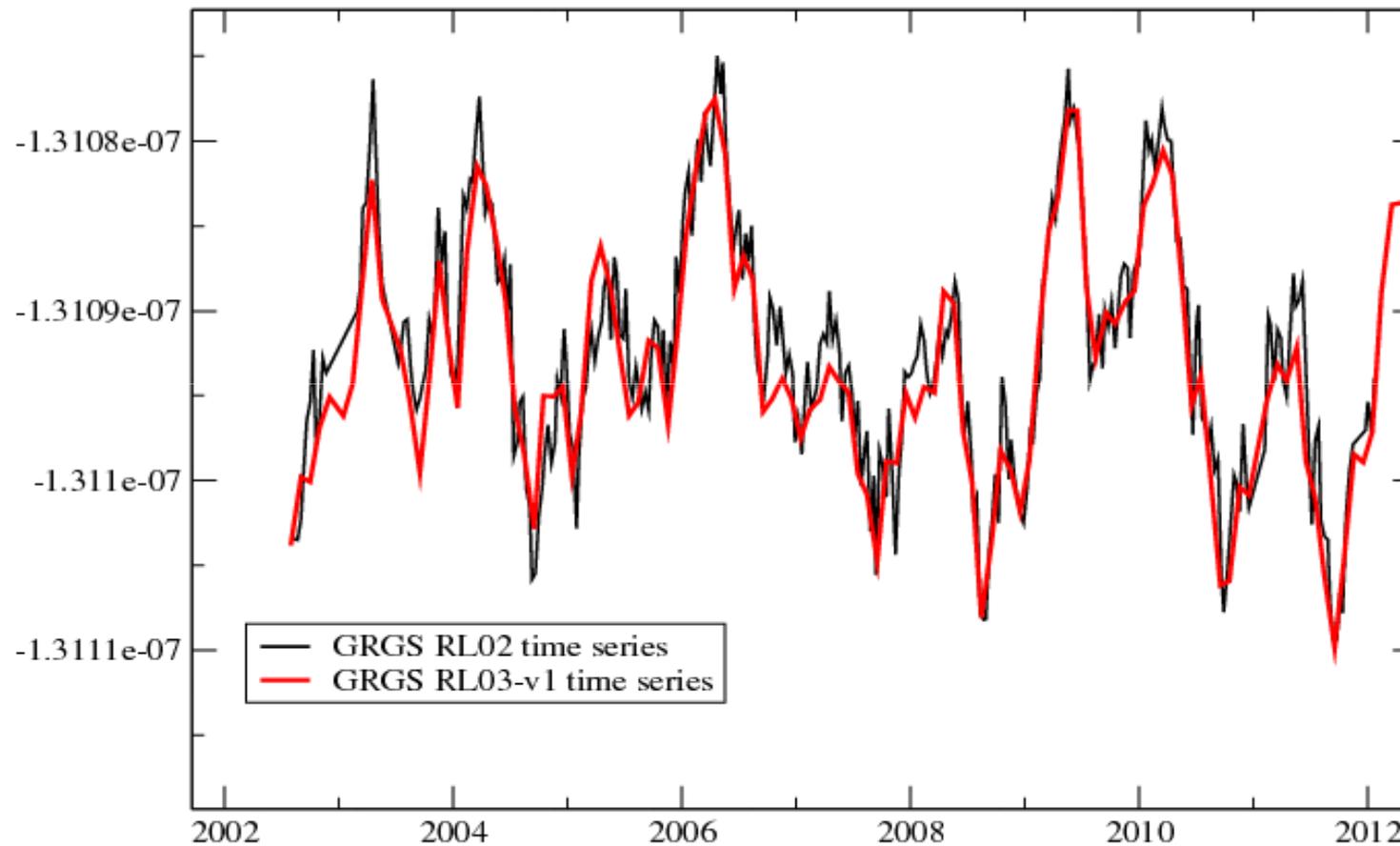


Reference uncertainty (qsum = 0.87 cm)

## ❖ Mean models are now generated from time series

- Fitting each series of monthly spherical harmonic coefficients by a set of 6 parameters :
  - Yearly bias and slope : piecewise linear function except in case of ...
  - Jumps caused by big earthquakes (3 so far : Sumatra, Concepcion and Tohoku)
  - Annual and semi-annual sine/cosine functions (with continuity constraints at hinge epochs)
- It means 600 000 coefficients for a 80x80 spherical harmonic model which better match with GRACE monthly models
- Used for operational computation (i.e. altimetric orbit processing) or TRF processing (i.e. ITRF2014)

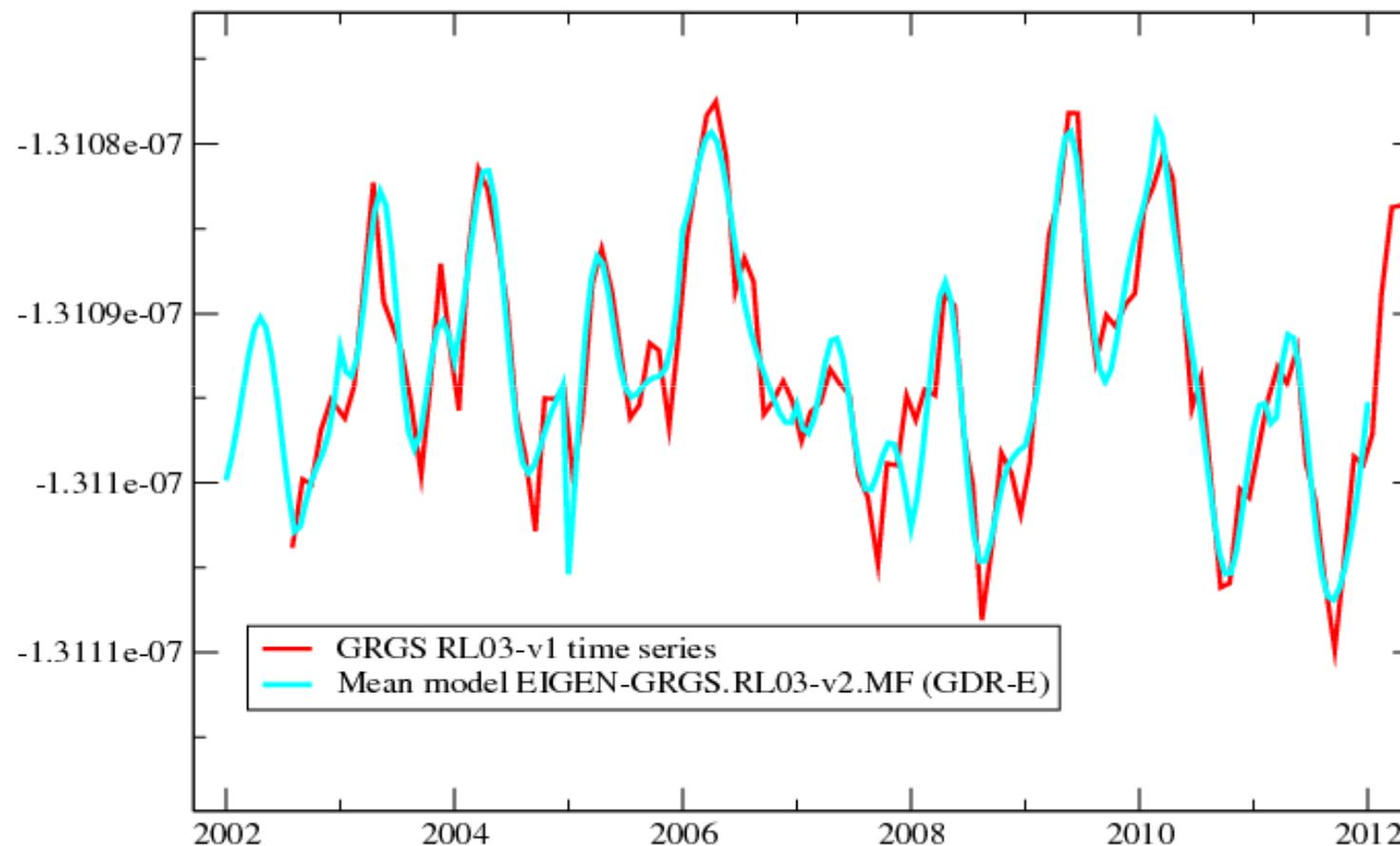
Normalized S (10,01) coefficient



In black: from RL02  
10-day series

In red : from RL03  
monthly series

Normalized S (10,01) coefficient

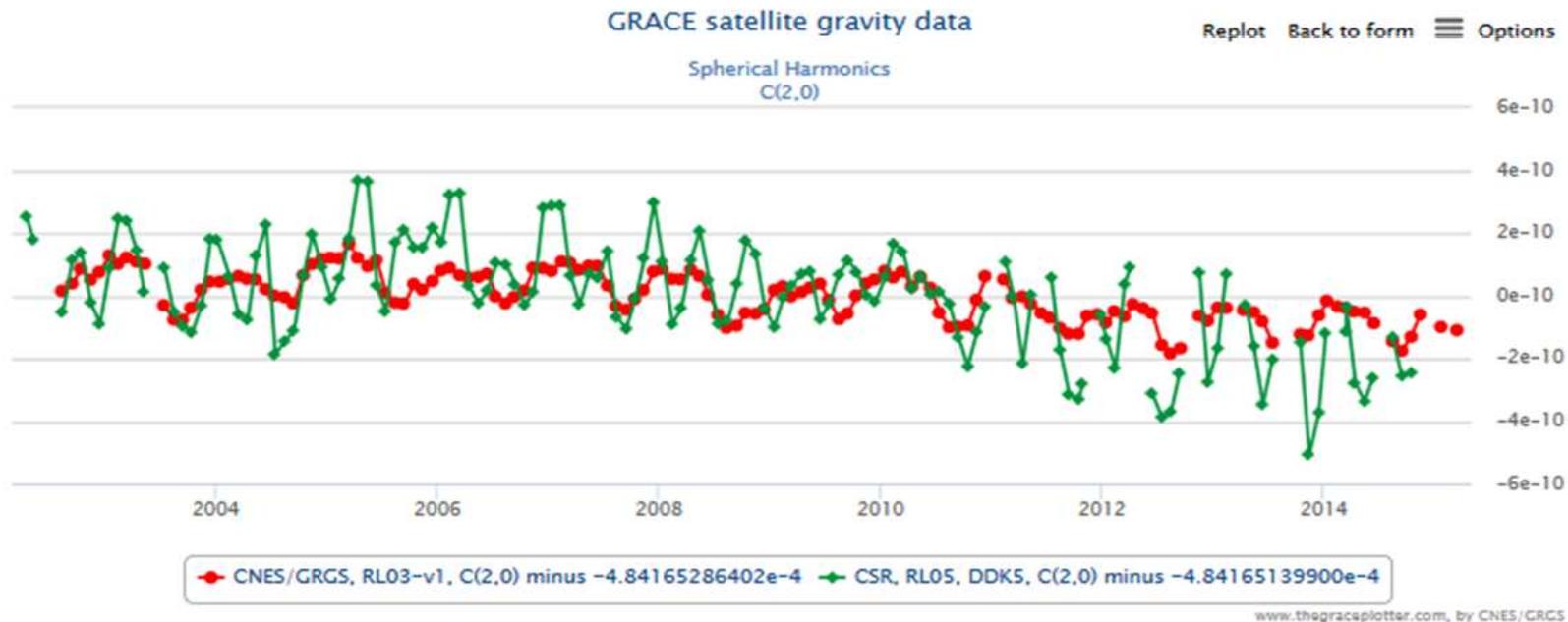


From RL03-v2 mean model with bias, drift per year, annual and semi-annual periodic terms per year

## ❖ J2 monthly variations are extended from 1986 till now

- From Lageos, Starlette and Stella data
- Need to be consistent with other time variable models, e.g. ocean tides ( $S_{sa}$ ,  $S_a$ ,  $\Omega_1$ )

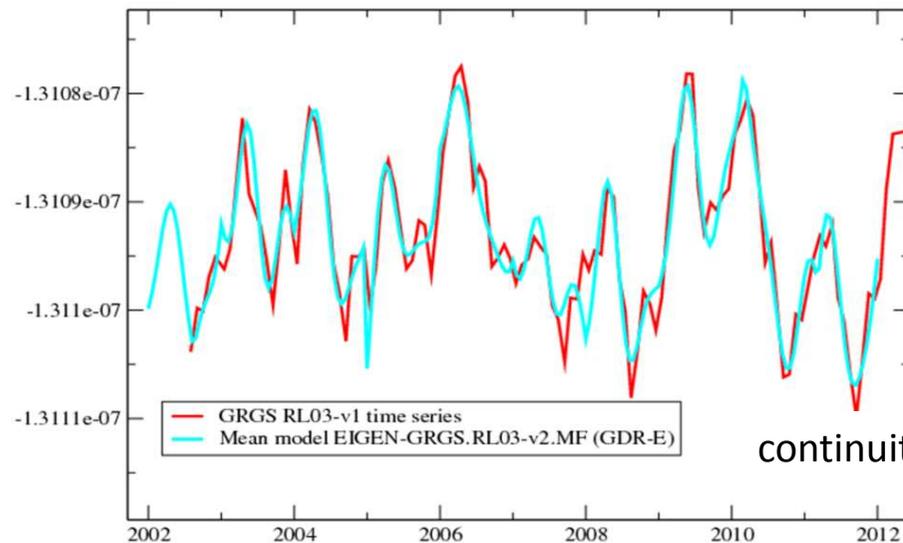
## ❖ [www.thegraceplotter.com](http://www.thegraceplotter.com) (CNES/GRGS)



## ❖ Extrapolated coefficients

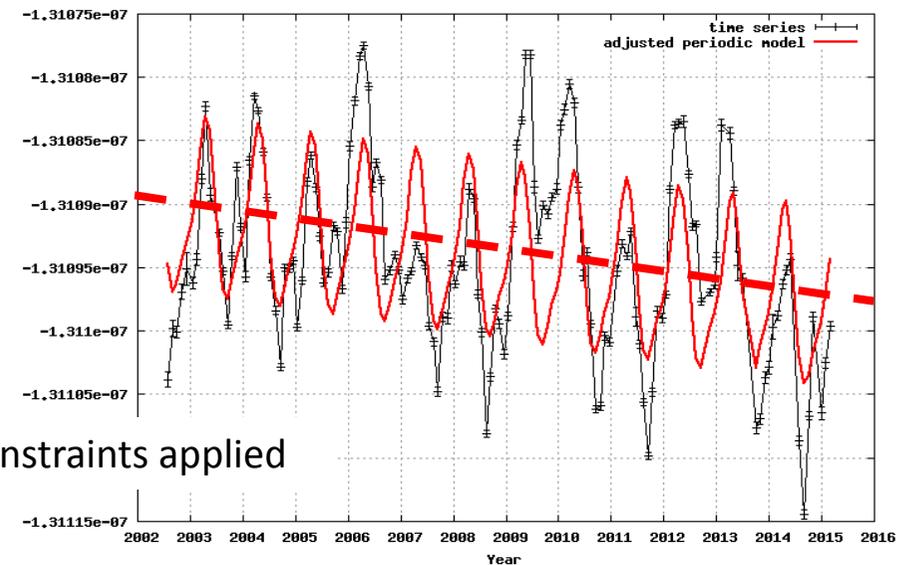
- Mean drift, mean annual and semi-annual periodic terms from the first (backward) and last (forward) determined biases :
  - Before 1986 for 2-degree terms determined from Lageos data
  - Before August 2002 for all other terms up to degree/order 80
  - From 2014 forward for all terms (RL03-v2 model)

S(10,1) within the GRACE period



continuity constraints applied

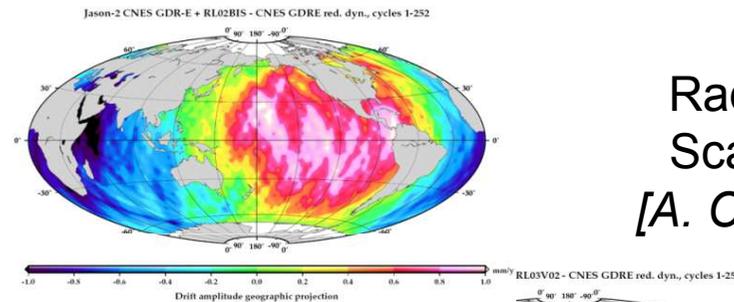
S(10,1) outside the GRACE period



- ❖ **The new RL03-v2 model** reduces the geographically correlated radial orbit drift rate, from more than 1 mm/yr (for the RL02bis mean model) to less than 0.6 mm/y over  $\sim 7$  years, with respect to Jason-2 GDR-E reduced-dynamic orbits (from GPS+DORIS).

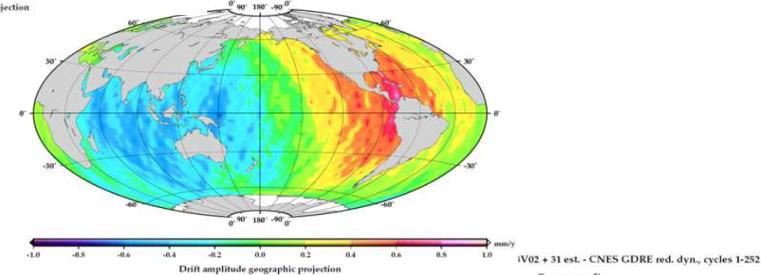
- ❖ **Jason-2 SLR residuals :**

- RL02: 1.36 cm rms

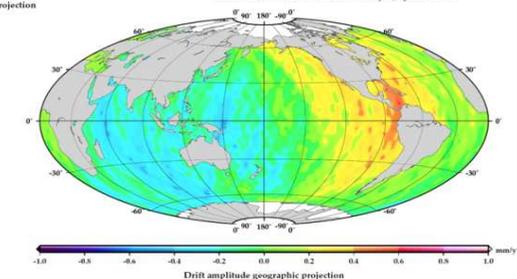


Radial orbit drift rate  
Scale: -1 / +1 mm/yr  
[A. Couhert & al., 2015]

- RL03-v2: 1.29 cm rms



- RL03-v2 + C31 adjusted: 1.27 cm rms



- RL02-v2 is used for GDR-E orbit production
- It is expanded in sets of 6 yearly coefficients (bias, slope, annual and semi-annual terms) per degree/order up to 80, and contains constant terms (from EIGEN6-S2) up to degree/order 175
- Extrapolated time-variable terms (before August 2002 and after July 2014) are based on global fits of monthly coefficients over 12 years of GRACE data
- Degree 2 time-variable terms are adjusted back to 1986 from Lageos data
- RL02-v2 is available on: [http://grgs.obs-mip.fr/grace/variable-models-grace-lageos/mean\\_fields](http://grgs.obs-mip.fr/grace/variable-models-grace-lageos/mean_fields)

## ❖ Next RL03-v3 model

- Improving the inversion process (Cholesky + SVD in a 2-step procedure)
- Completing the years 2014-2015 (no longer extrapolated)
- Homogenizing the relative weights (between GPS and KBR)
- Using more satellite data (Starlette, Stella, Lares, Jason...)

